CHARACTERIZATION OF FLOW DYNAMICS IN COASTAL AQUIFERS BASED ON HYDRAULIC TOMOGRAPHY.

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Submarine groundwater is richer then surface water, and has been considered to be an essential component of biological production in marine coastal ecosystems. (Fujita et al., 2019).

Submarine groundwater discharge (SGD) links land and marine systems, but has often been overlooked in coastal nutrient budgets because it is difficult to quantify (R. Santos et al., 2021).



■ Influenced by oceanic oscillation and inland head → Complex flow and transport process (Li et al., 2010).





 Combine hydrogeological and geophysical data allowed for the construction of a hydrogeophysical model of the multi-layered system(R. Di Maio, 2014).

 In some cases, core, slug, geophysical data may be collected and use to condition the inverse modeling results(A. Illman, 2007).





- Determining aquifer parameters (K) by field measurement and inversion.
- Using the inverse method to estimate the hydraulic conductivity distribution.
- Using field experiment data with different scales to compare the differences in inverse results.





Σ	Introduction	Meth	nodology	Results & Di	iscussion	Future work	
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Well Number	Falling Head Test	Slug Test	Single Well Pumping Test	Multi-Well Pumping Test	Cross-hole Layered Pumping Test
BW01	\checkmark		\checkmark		
BW02			\checkmark	\checkmark	
BW03			\checkmark	\checkmark	\checkmark
BW09		\checkmark	\checkmark	\checkmark	\checkmark
BW10		\checkmark	\checkmark	\checkmark	\checkmark
BW11	\checkmark			\checkmark	\checkmark





- Auto monitoring: 1 data/hour, 1 month, 5 wells.
- Purpose: set the boundary of the model.
- Due to the tidal period, the equipment at least set up for 1 month.











- Aquifer has infinite areal extent
- Aquifer is homogeneous, anisotropic and of uniform thickness
- Control well is fully or partially penetrating
- Diameter of a pumping well well is very small so that storage in the well can be neglected

Analysed by Theis solution.









Separate different layers to do the slug test : every 2 meters.







- Determine K Value.
- Choose a specific depth to do the experiment.







- VSAFT2 is a windows Graphical User Interface (GUI) for setting up, running and calibrating a variably saturated flow and transport finite element model (Yeh, et al., 1993) in two-dimensional.
- VSAFT2 now includes several geostatistical model setup features such as random field generation of input parameters.







VSAFT 2



	Mate	ial	Ksx	Ksy	
	Blac	k	1.57	1.57	
			Stress 1	Stress 2	
BW09			-48.096	OBS	
BW10			OBS -48.09		
BW11		/	OBS		
BW03			OBS		
Grid		-	35*30(row*col)		
Flow			Steady State		
Material			Heterogeneous		
Initial condition			Pressure head : 4.317m		
Well			4 wells		
Boundary			Left: pres.head/-4.317 Right: pres.head/4.513m		







• Finish all lab and field experiment.

• Modify the inverse K field.

• Using the K field to construct the hydraulic tomography.

THANK YOU FOR LISTENING

$$\xi_*^{(1)} = \lambda^T \zeta_{\text{obs}} + \mu^T \varepsilon_{\text{obs}}$$
$$\xi_*^{(k+1)} = \xi_*^{(k)} + \omega^{(k)} (\varepsilon_{\text{obs}} - \varepsilon^{(k)})$$

Ni, et al., (2009)

 $\xi_*^{(k)}$ is the kth estimation of parameter (i.e., the hydraulic* conductivity K), ζobs and εobs represent the differ-ences of the parameters (i.e., hydraulic conductivity) and hydraulic heads at observation locations,

 $\lambda, \mu, \omega(k)$ are cokriging weighting vectors evaluated by stochastic simulations.



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- b is aquifer thickness [L]
- Kr is the radial (horizontal) hydraulic conductivity [L/T]
- Kz is the vertical hydraulic conductivity [L/T]
- ID is dimensionless depth to bottom of pumping well screen (I/b)
- Q is pumping rate [L³/T]
- r is radial distance from pumping well to observation well [L]
- s is drawdown [L]
- S is elastic storage coefficient [dimensionless]
- Sy is specific yield [dimensionless]
- ts is dimensionless time with respect to S
- t is elapsed time since start of pumping [T]
- T is transmissivity [L²/T]





Parameter	Value	Unit
а	7.07	cm ²
L	15	cm
А	44.16	cm ²
ho	84	cm
h1	54	cm
log10(ho/h1)	0.19	

$$\mathsf{K}=2.3\frac{aL}{A(t_{1}-t_{0})}\log_{10}\frac{h_{0}}{h_{1}}$$

-a: the area of standpipe
-L: the length of sample
-A: the area of sample
-t₁-t₀: the time for falling from h₀ to h₁



