

# **APPLYING THE VARIABLY SATURATED FLOW MODEL TO SIMULATE GROUNDWATER FLOW IN PINGTUNG PLAIN BY USING THMC**

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**Date:** 20<sup>th</sup> Oct 2023

# Outline

I INTRODUCTION

II METHODOLOGY

III PRELIMINARY RESULT

IV FUTURE WORK

# I INTRODUCTION

II

III

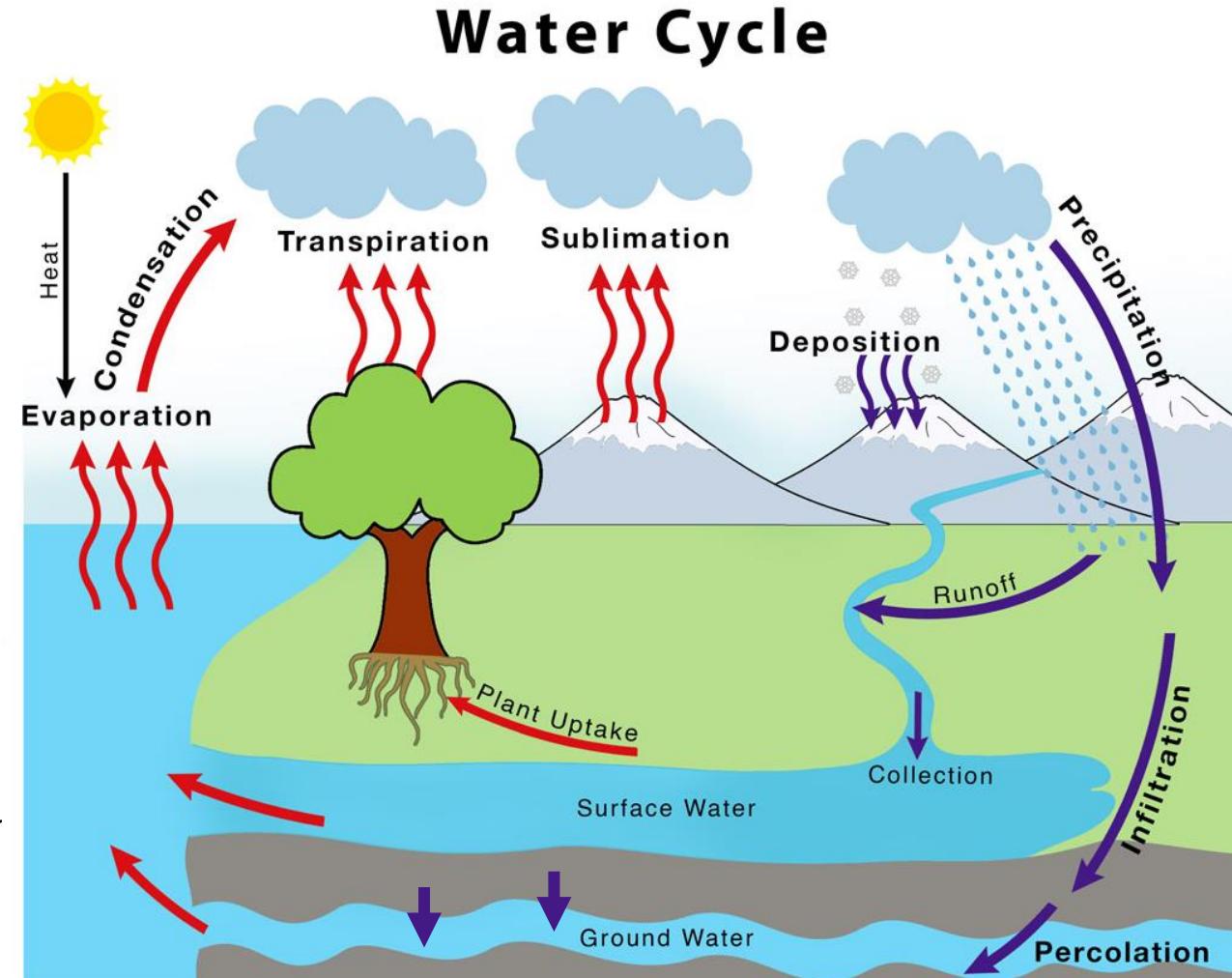
IV

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Groundwater is an important water resource for human life with different purpose of using:



The groundwater withdrawal amount in Asia account for the majority is 72% of global usage. (Lee et al., 2018)



Water cycle (sciencefacts)

# I INTRODUCTION

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IV

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In Taiwan, tap water coverage average rate is 92.7 %

Limited surface water

Lowest ratio of tap water use (45.8%)

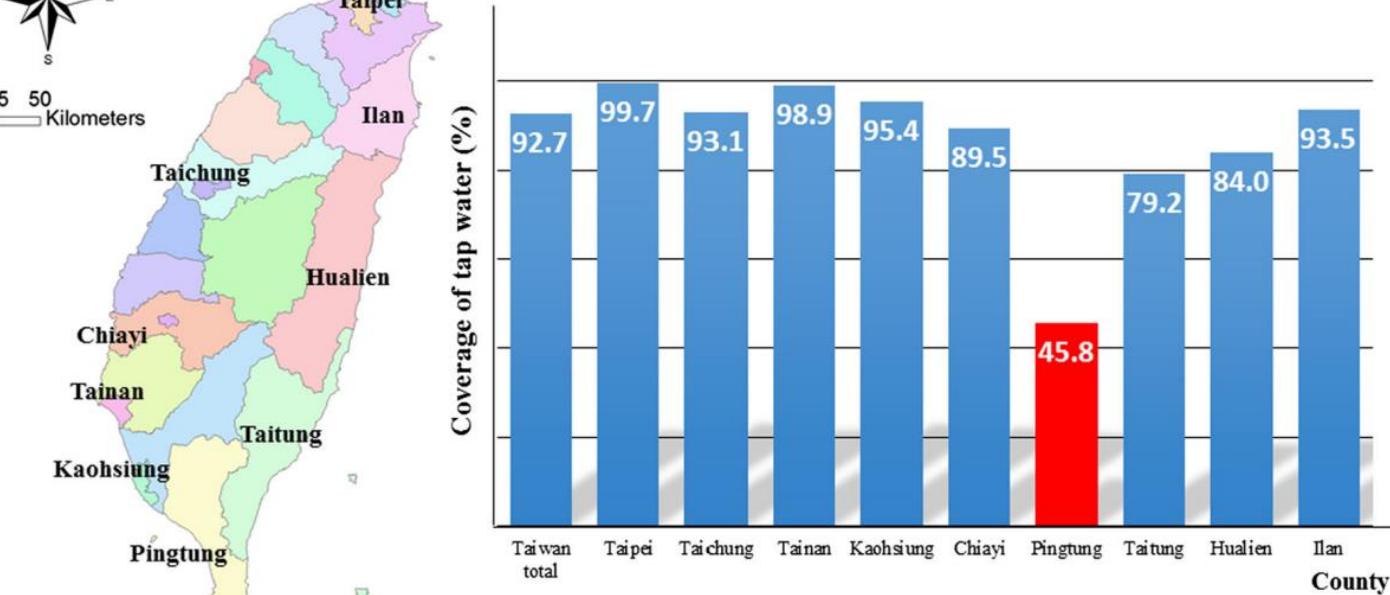
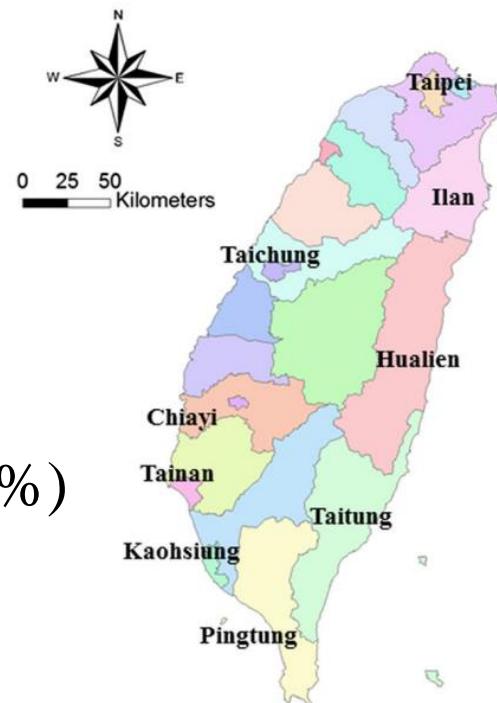
Over-exploitation groundwater



Seawater intrusion, and land subsidence in over the past 5 decades (Ting et al., 1998)



Understanding the behavior of groundwater – groundwater flow with the aim of using this resource in a reasonable and sustainable way.



*The coverage of the tap water in main city in Taiwan (Liang et al., 2016)*

# I INTRODUCTION

## II

## III

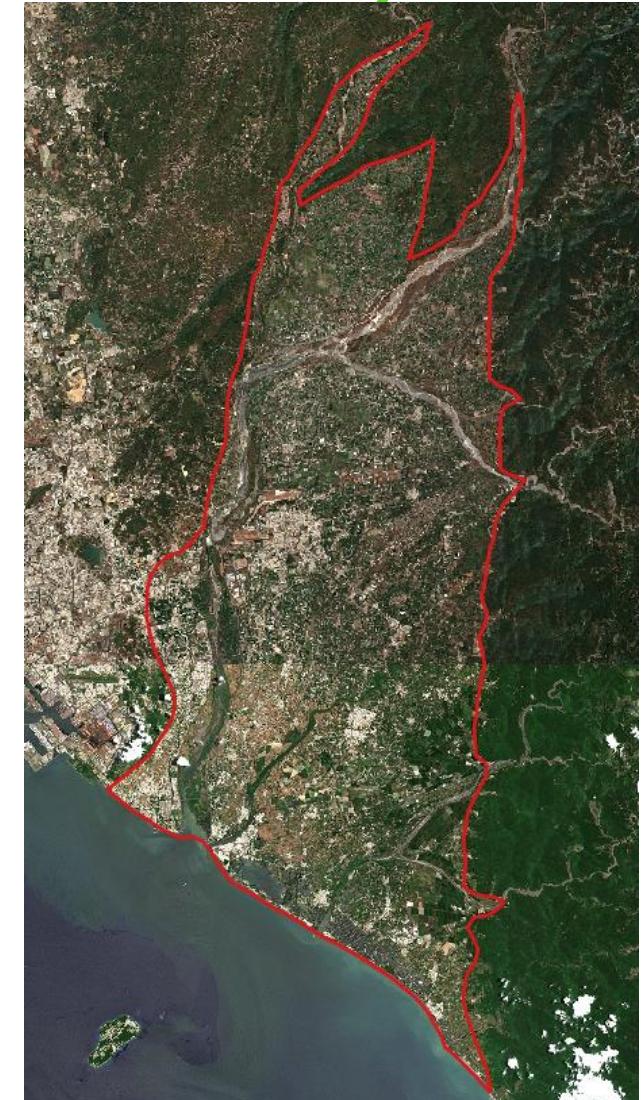
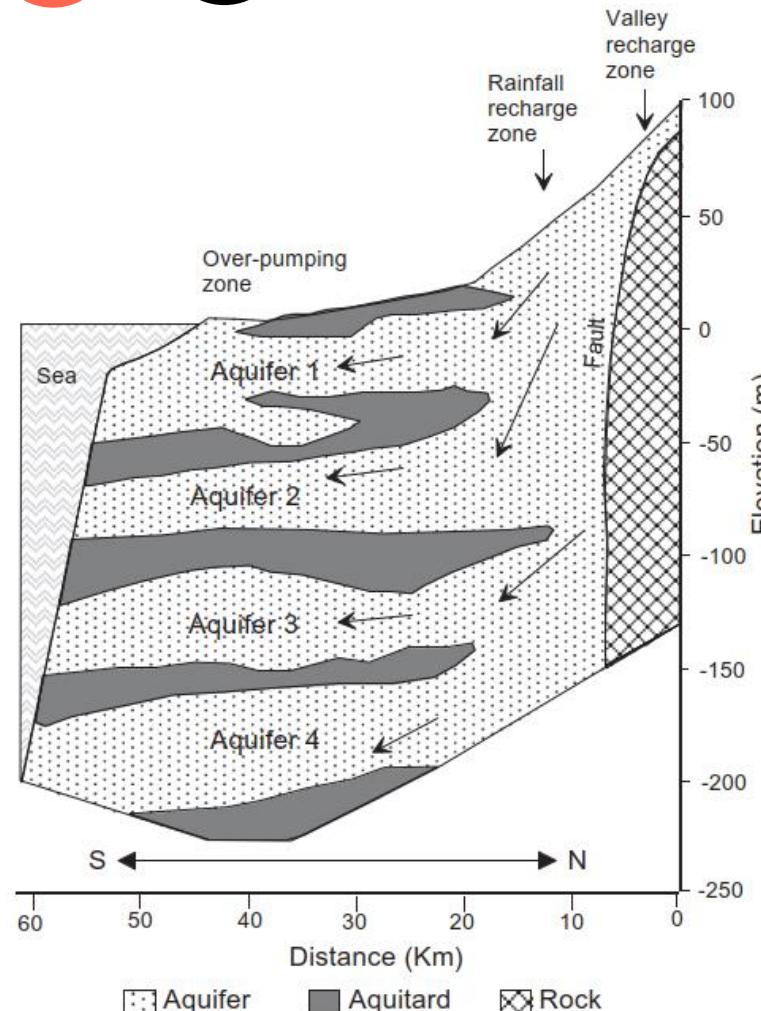
## IV

### Study area - Pingtung plain

- 1210 km<sup>2</sup>
- Natural boundary
- High annual rainfall rate
- Unconsolidated Quaternary

### Hydrogeological

- Aquifer: high permeable coarse sediment.
- Aquitard: low permeable fine sediments



### Objective

Applying the variably saturated flow model to simulate groundwater flow in Pingtung plain by using THMC

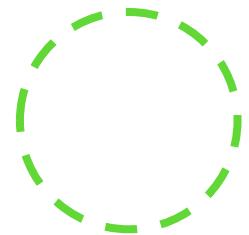
# I INTRODUCTION

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## THMC evolution (Thermal – Hydrology – Mechanic – Chemical)



### HYDROGEOCHEM

- ✓ Water flow
- ✓ Water quality transmission & chemical reaction
- ✓ Heat transfer

1981

1991

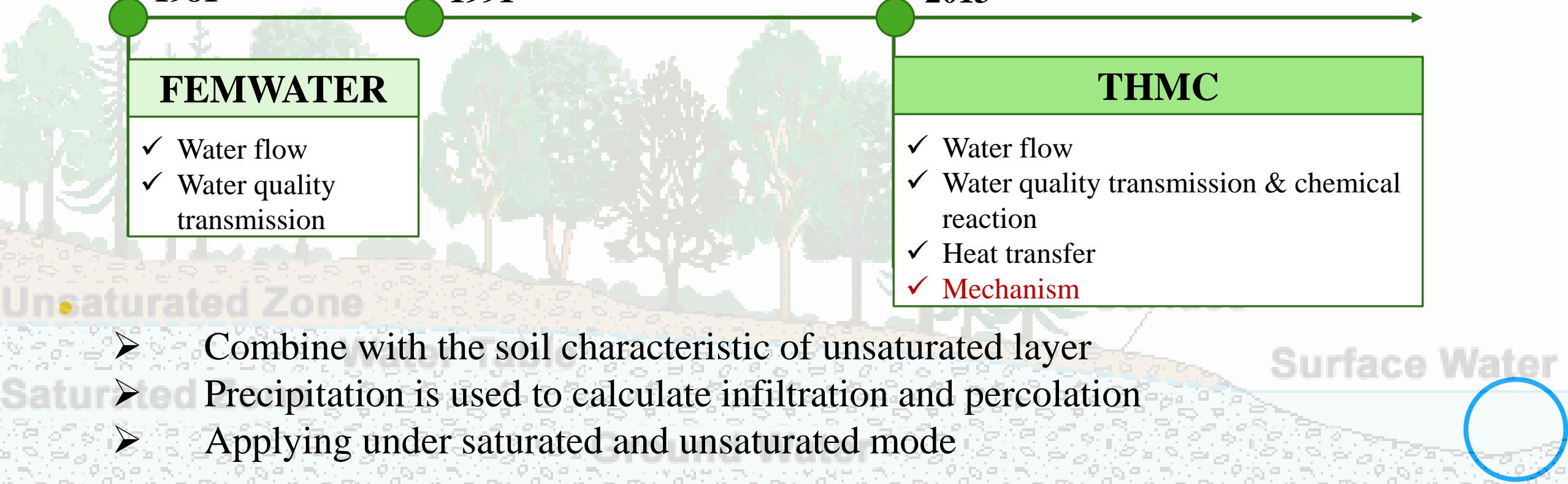
2013

### FEMWATER

- ✓ Water flow
- ✓ Water quality transmission

### THMC

- ✓ Water flow
- ✓ Water quality transmission & chemical reaction
- ✓ Heat transfer
- ✓ Mechanism



# I INTRODUCTION

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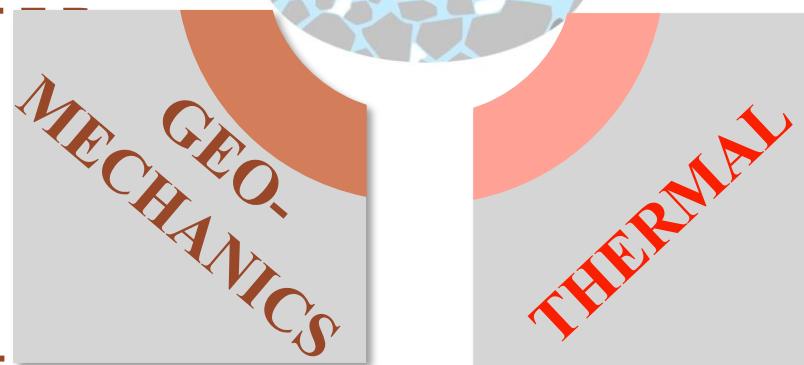
## THMC software application

- ✓ Wide application
- ✓ Graphical Users Interface (GUI) display
- ✓ Accurately predict subsurface process



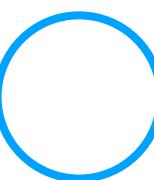
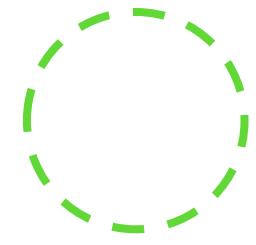
- Groundwater management
- Groundwater recharger

- Water quality assessment
- Seawater intrusion
- Land utilization management



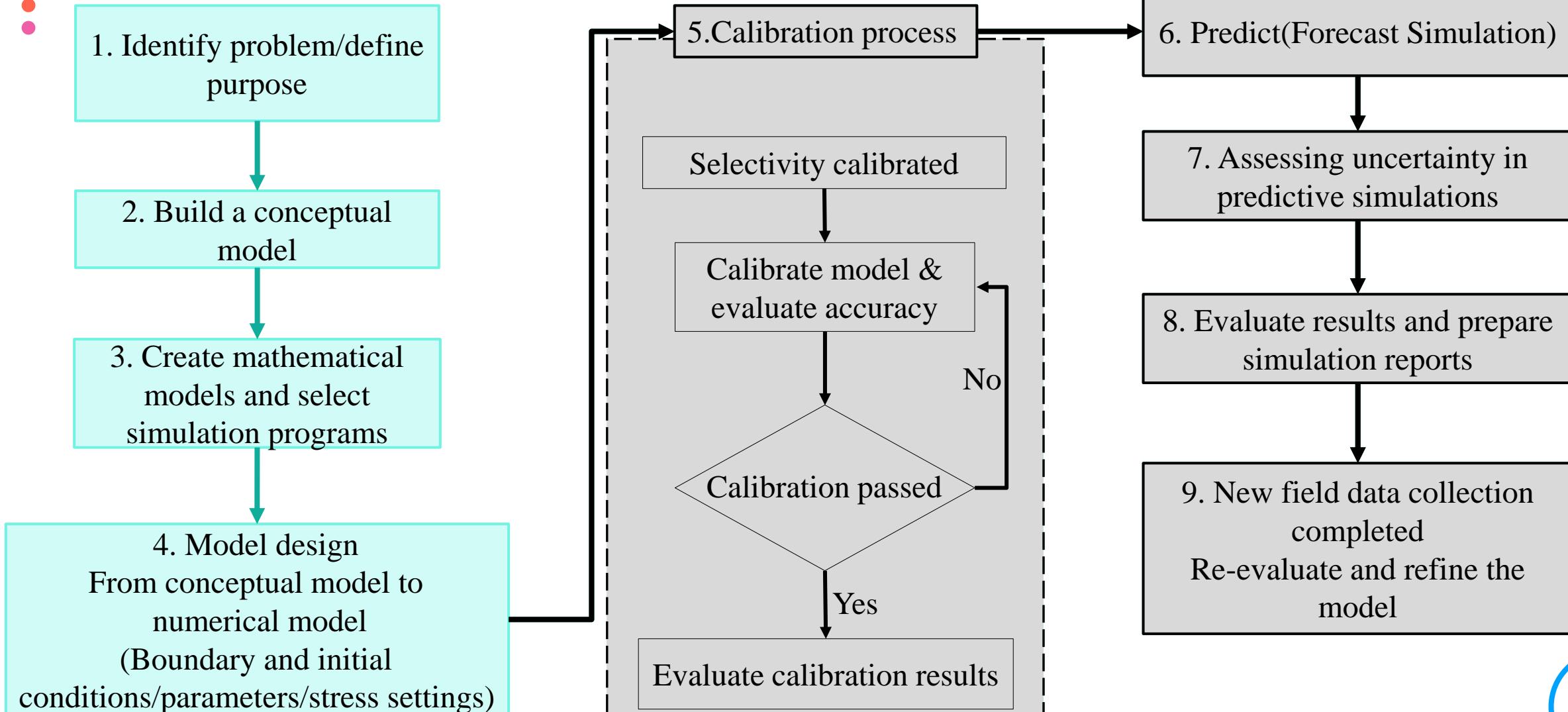
- Slope area management
- Subsidence assessment
- Conservation of soil and water assessment

- Hot spring groundwater
- Geothermal energy



## **II METHODOLOGY**

## Working flow



## Governing equation

## Boundary Condition

## Data input

Governing equation for flow through saturated-unsaturated media using in software follow below equation:

$$\frac{\rho}{\rho_0} F \frac{\partial h}{\partial t} = \nabla \cdot \left[ K \cdot \left( \nabla h + \frac{\rho}{\rho_0} \nabla z \right) \right] + \frac{\rho^*}{\rho_0} q$$

(Yeh et al., 1994a, 1994b)

$\theta$ : effective moisture content ( $L^3/L^3$ )

$h$ : pressure head (L)

$t$ : time (T)

$z$ : potential head (L)

$q$ : source/sink of fluid [ $(L^3/L^3)/T$ ]

$\rho_0$ : referenced fluid density at zero chemical concentration ( $M/L^3$ )

$\rho$ : fluid density with dissolved chemical concentrations ( $M/L^3$ )

$\rho^*$ : fluid density of either injection ( $\rho^*$ ) or withdraw (=  $\rho$ )

$\mu_0$ : fluid dynamic viscosity at zero chemical concentration [ $(M/L)/T$ ]

$\mu$ : fluid dynamic viscosity with dissolved chemical concentrations [ $(M/L)/T$ ]

$\alpha'$ : modified compressibility of the soil matrix (1/L)

$\beta'$ : modified compressibility of the liquid (1/L)

$n_e$ : effective porosity ( $L^3/L^3$ )

$S$ : degree of effective saturation of water

$g$ : gravity ( $L/T^2$ )

$k$ : permeability tensor ( $L^2$ )

$k_s$ : saturated permeability tensor ( $L^2$ )

$K_{so}$ : referenced saturated hydraulic conductivity tensor (L/T)

$k_r$ : relative permeability or relative hydraulic conductivity (dimensionless)

$F$ : generalized storage coefficient (1/L)

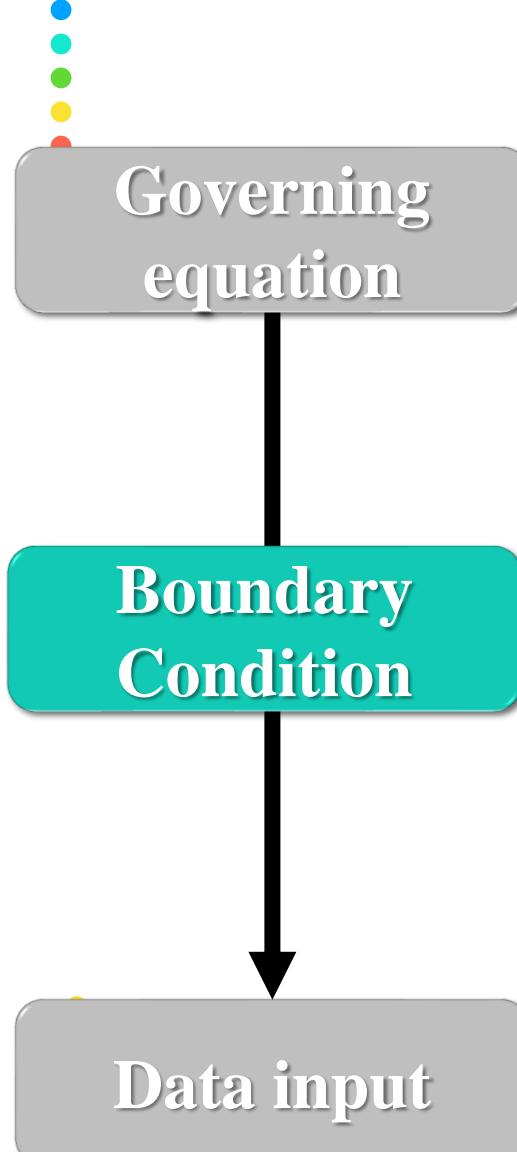
$$F = \alpha' \frac{\theta}{n_e} + \beta' \theta + n_e \frac{dS}{dh}$$

$K$ : hydraulic conductivity tensor (L/T)

$$K = \frac{\rho g}{\mu} k = \frac{\rho}{\mu} \frac{\rho_0}{\mu_0} \frac{\rho_0 g}{\mu_0} k_s k_r = \frac{\rho}{\mu} \frac{\rho_0}{\mu_0} K_{so} k_r$$

Darcy's velocity (L/T)

$$V = -K \cdot \left( \frac{\rho_0}{\rho} \nabla h + \nabla z \right)$$

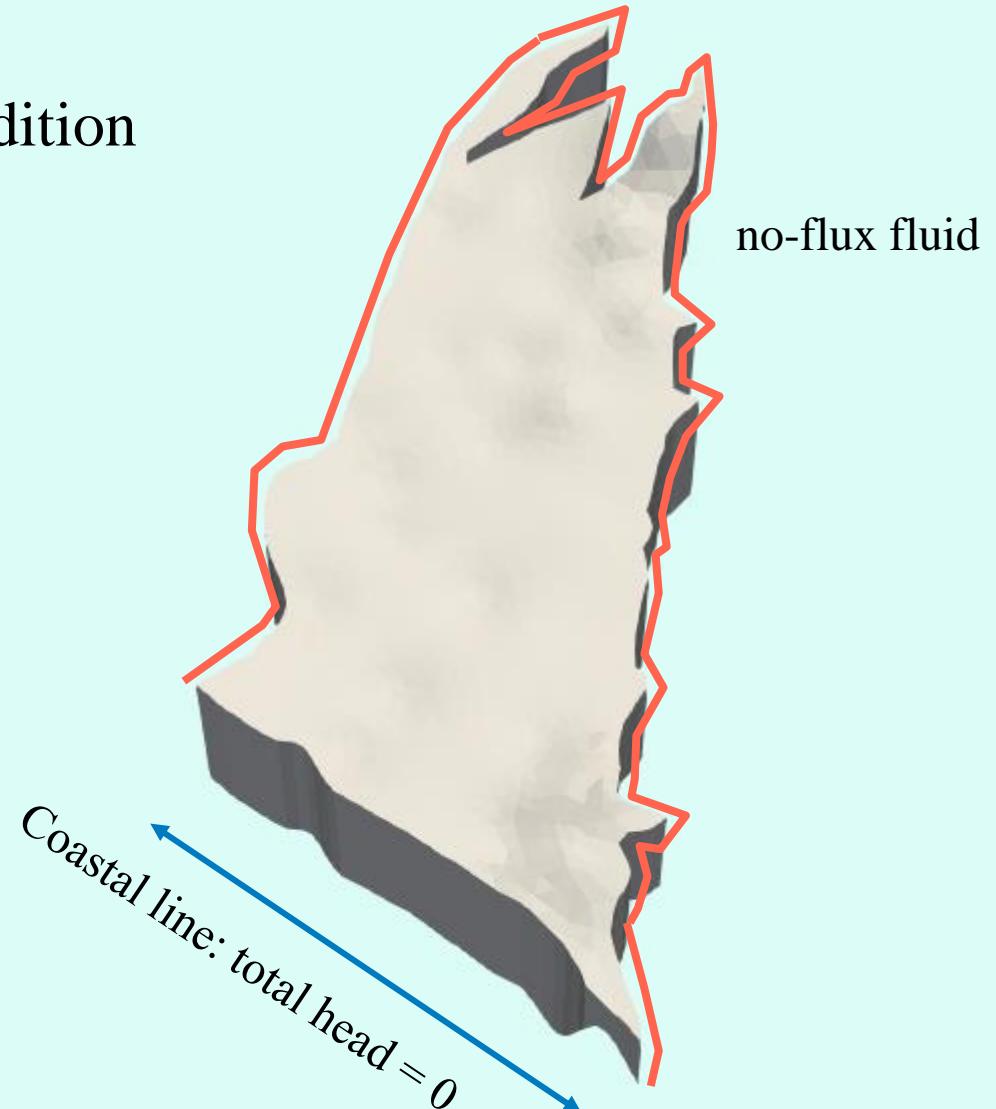


## Steady-state simulation for Initial condition

- ✓ Dirichlet boundary condition
- ✓ Rainfall: variable BC

Initial result

Transient simulation

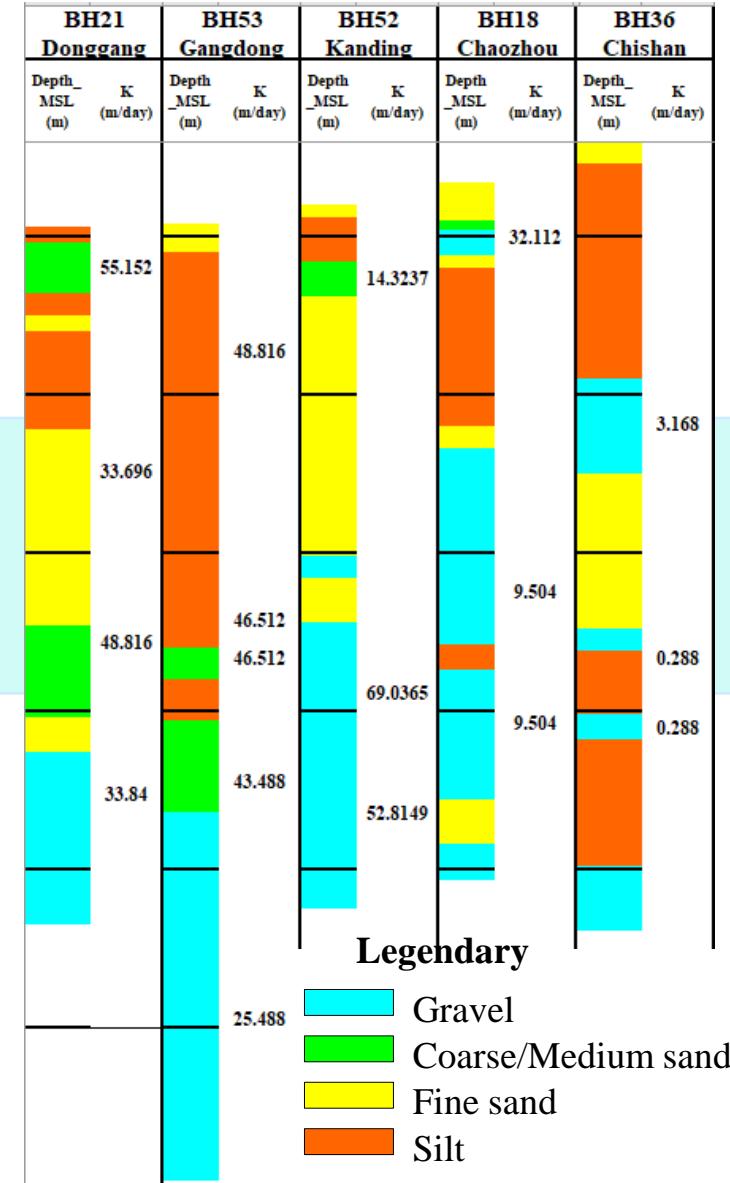
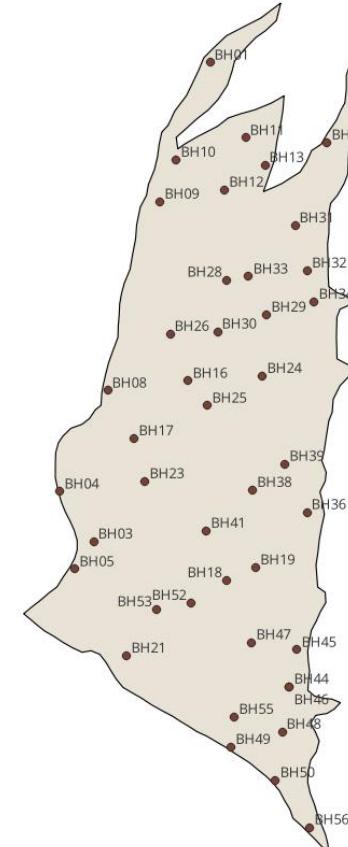


## Governing equation

## Boundary Condition

## Data input

- Conceptual hydrological: build up from 42 borehole cores (CGS website)



1 material (1<sup>st</sup> layer)  
4 aquifers  
3 aquitards

## Governing equation

## Boundary Condition

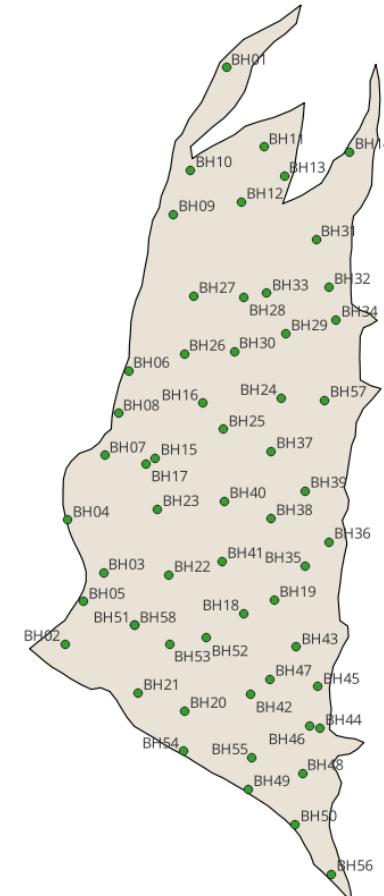
## Data input

- K value reference:

Aquifer: 148 wells → Pumping test result from CGS report

Aquitard: According to type of rock and giving K value base on reference table of Domenico and Schwartz (1998)

Layer	Material	K value range (m/day)	Average K (m/day)
1	Material 8	0.8640 – 8.64864	5.7294
2	Aquifer1	0.0057 – 461.52	89.681
3	Aquitard1	0.8640 – 8.64864	4.1285
4	Aquifer2	0.4320 – 171.792	34.291
5	Aquitard2	0.8640 – 8.64864	2.9126
6	Aquifer3	2.2920 – 171.936	40.076
7	Aquitard3	0.8640 – 8.64864	6.0537
8	Aquifer4	5.472 – 14.5152	9.5688



Governing  
equation

Boundary  
Condition

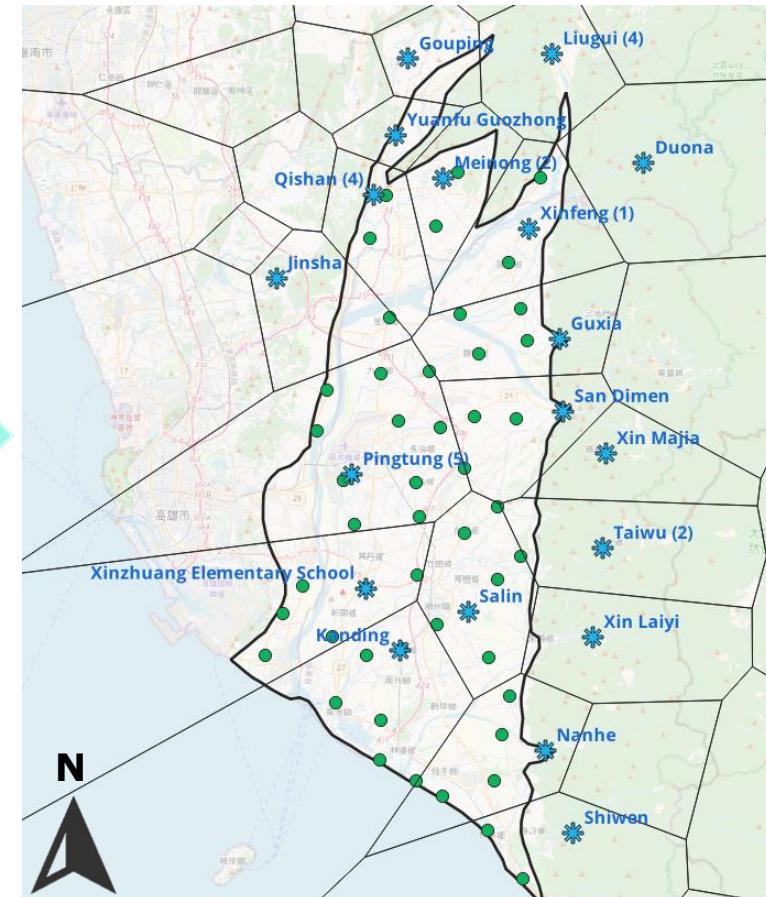
Data input

Groundwater level: daily data 2020 – 2022 of 122 stations (WRA)  
Rainfall rate: daily data 2020 – 2022 of 19 stations (WRA)

Station ID	Station name	2020	2021	2022	Unit: mm/year
1670P001	Jinsha	2192	3255	1244	
1730P021	Guxia	2286	3698	1350	
1730P060	Pingtung (5)	1996	3536	1241	
1730P081	Meinong (2)	1744	3276	1580	
1730P107	Qishan (4)	1770	3277	1380	
1730P100	San Dimen	2102	3102	1462	

Using Thiessen method to determine  
rainfall distribution zone

1730P150	Gouping	1797	1437	
1730P151	Yuanfu Guozhong	1631	3106	1522
1730P152	Xinzhuang School	2118	3141	1077
1740P049	Salin	2136	3737	1413
1740P050	Kanding	1941	3447	1182
1740P051	Taiwu (2)	3270	5658	1762
1760P011	Nanhe	2123	3453	1117
1760P013	Xin Laiyi	2694	4962	1447
1790P002	Shiwen	2439	3521	1161



- GW level station
- ✿ Rainfall station

## Governing equation

## Boundary Condition

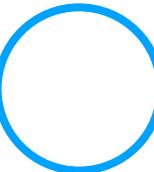
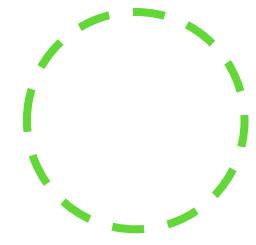
## Data input

Pumping rate: using groundwater usage right of Pingtung county (WRA) to represent for pumping rate

Unit:  $10^6 m^3$

Year	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec	Total
2020	77.89	70.40	78.00	75.27	77.64	71.99	77.58	77.12	74.62	77.29	75.14	75.34	908.26
2021	84.63	76.44	84.30	81.49	83.92	68.08	240.67	84.46	81.95	84.70	81.83	70.44	1122.92
2022	94.58	85.52	94.23	91.72	94.74	81.17	95.15	95.43	92.79	96.13	93.13	85.47	1100.06

Infiltration and percolation: using the land use/land cover to identify built area (building and road) which is assigned as impermeable surface.



### **III PRELIMINARY RESULT**

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II

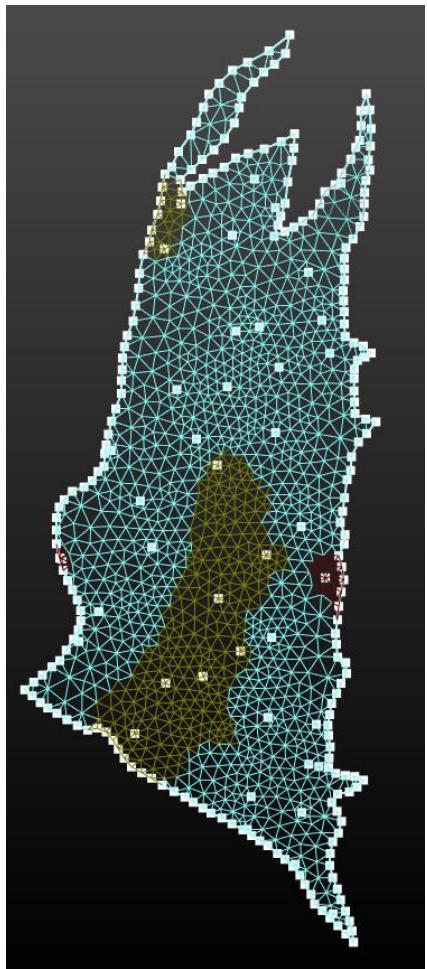
III

## PRELIMINARY RESULT

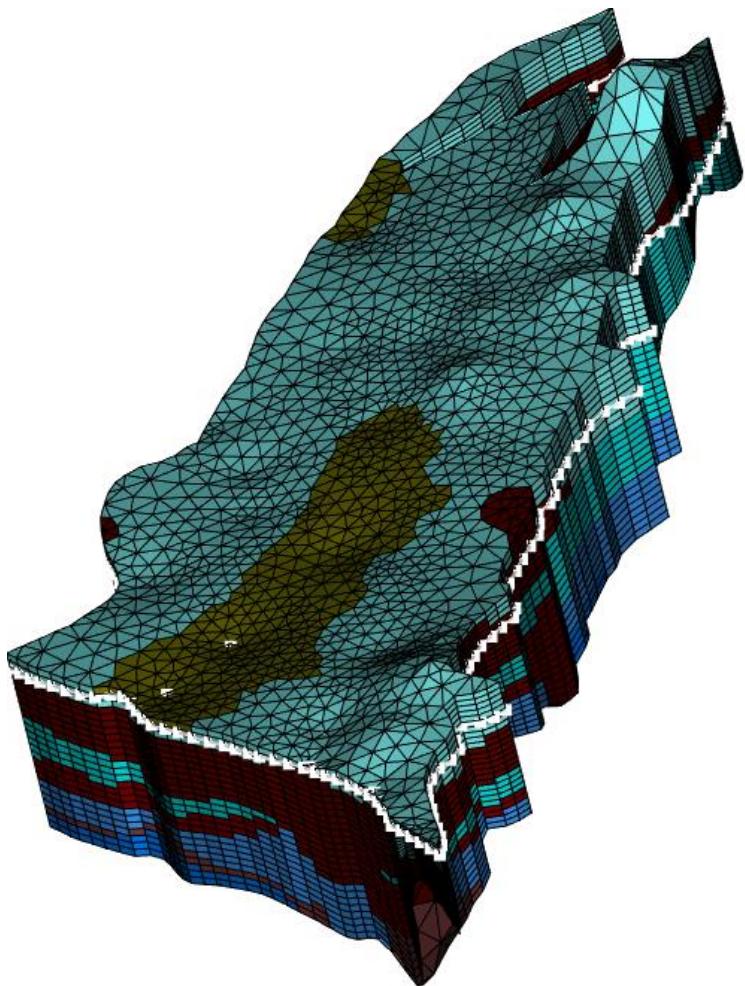
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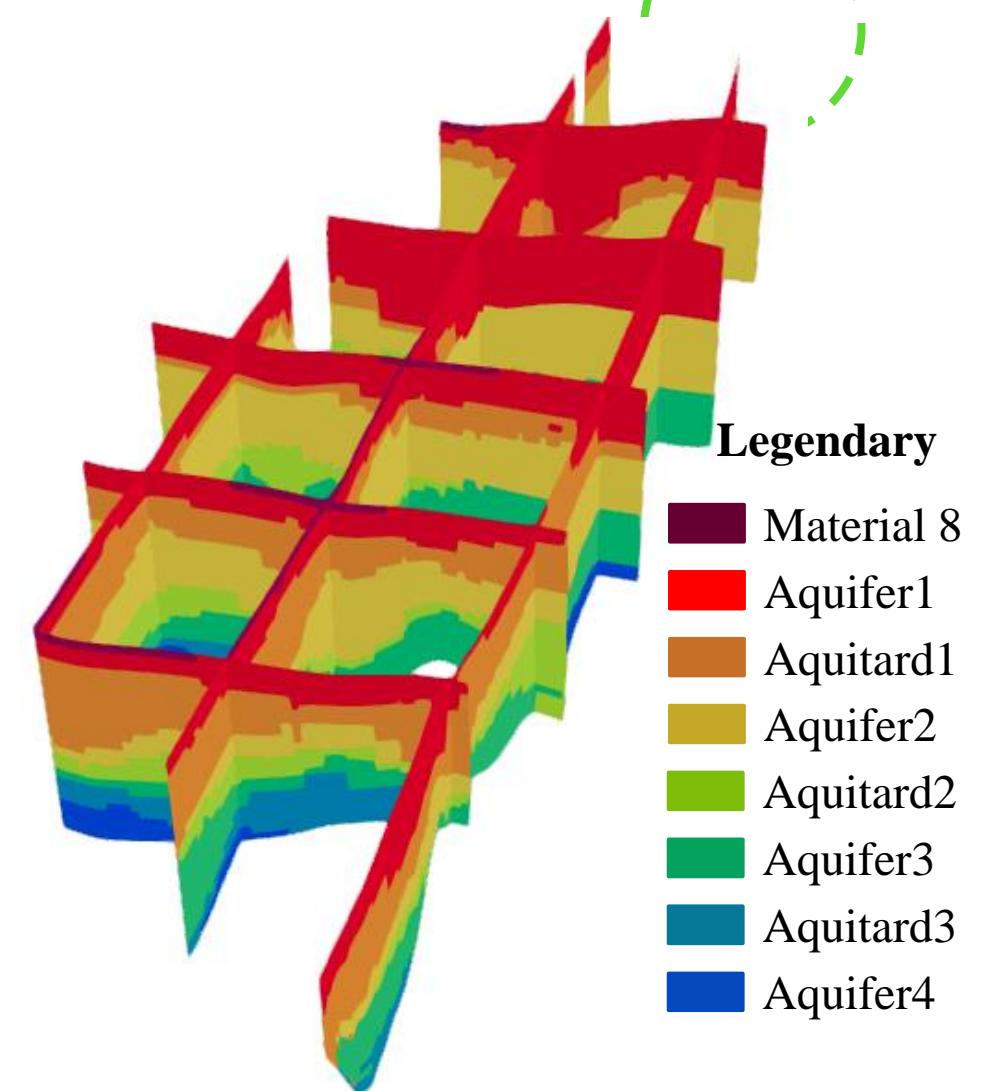
### Conceptual hydrological domain



Triangular column grid  
system with 33,826 nodes  
and 10,329 elements



3D Geological  
model generation



Cross-section of geological  
model

Legendary	
Material 8	
Aquifer1	
Aquitard1	
Aquifer2	
Aquitard2	
Aquifer3	
Aquitard3	
Aquifer4	

I

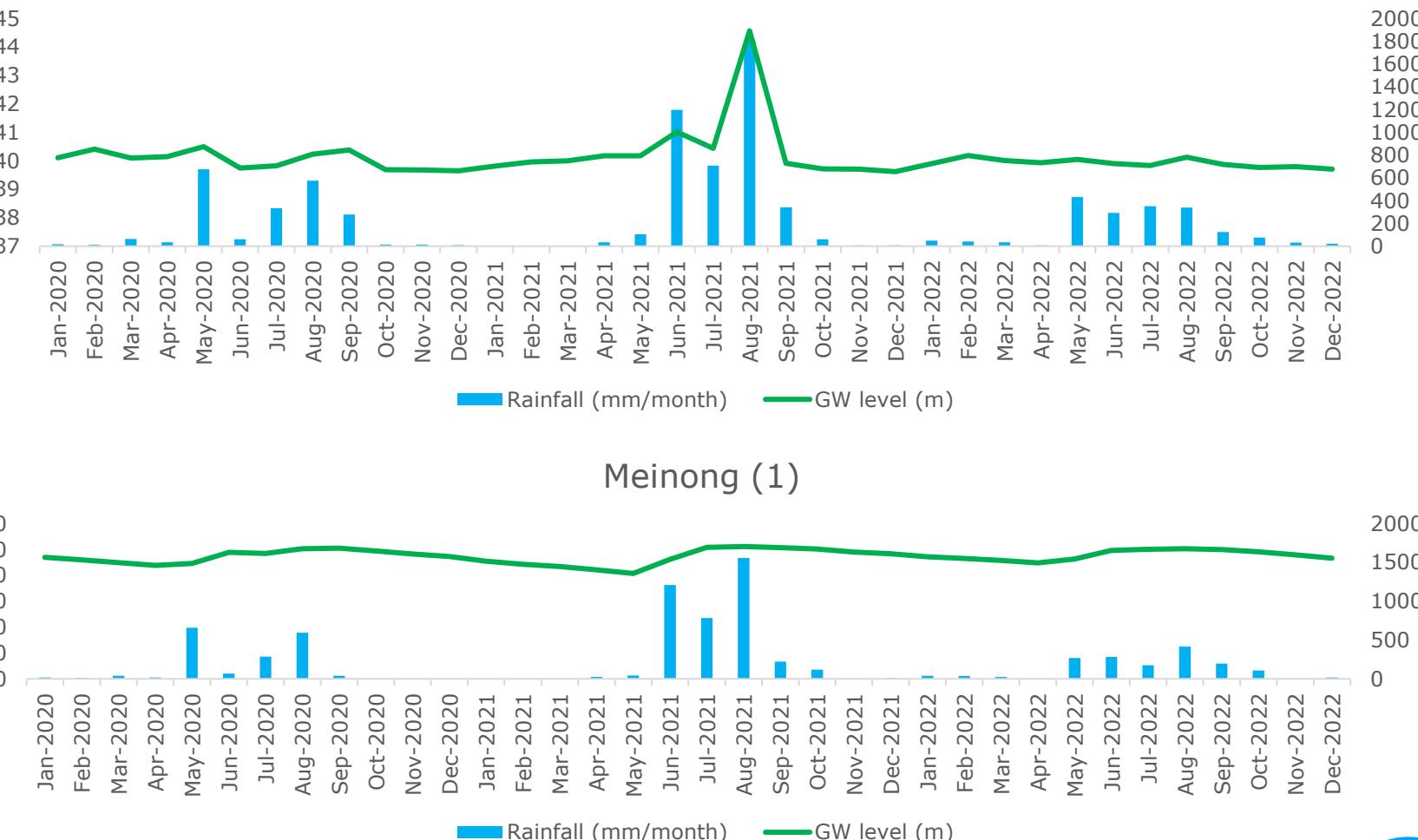
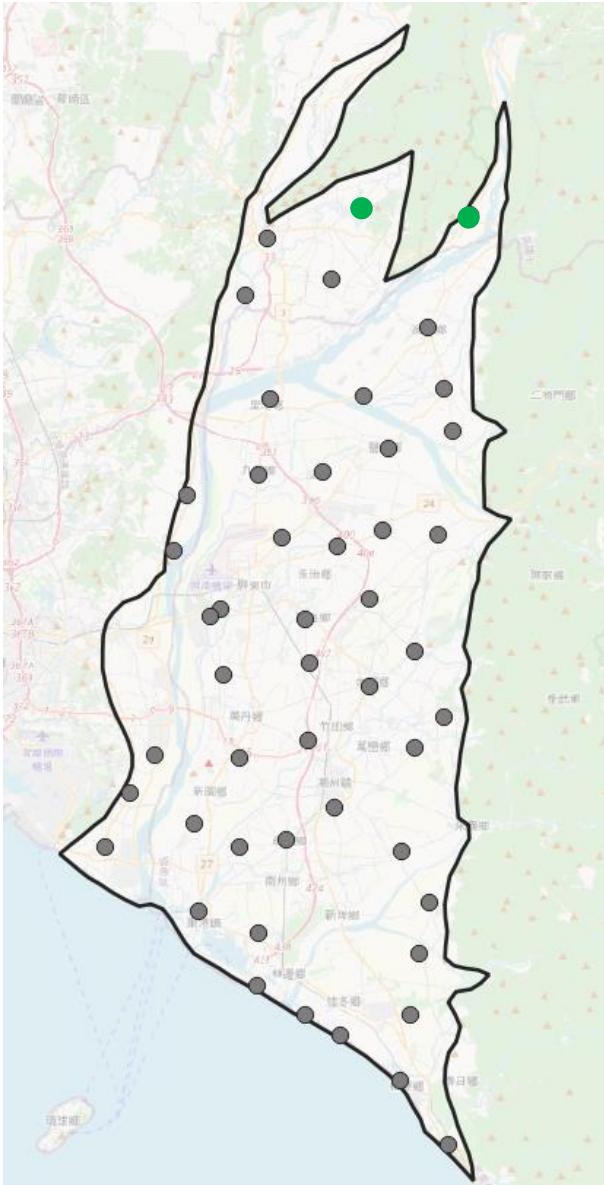
II

III

# PRELIMINARY RESULT

IV

## Groundwater level and rainfall distribution relationship



➤ Groundwater immediately increase after raining time

I

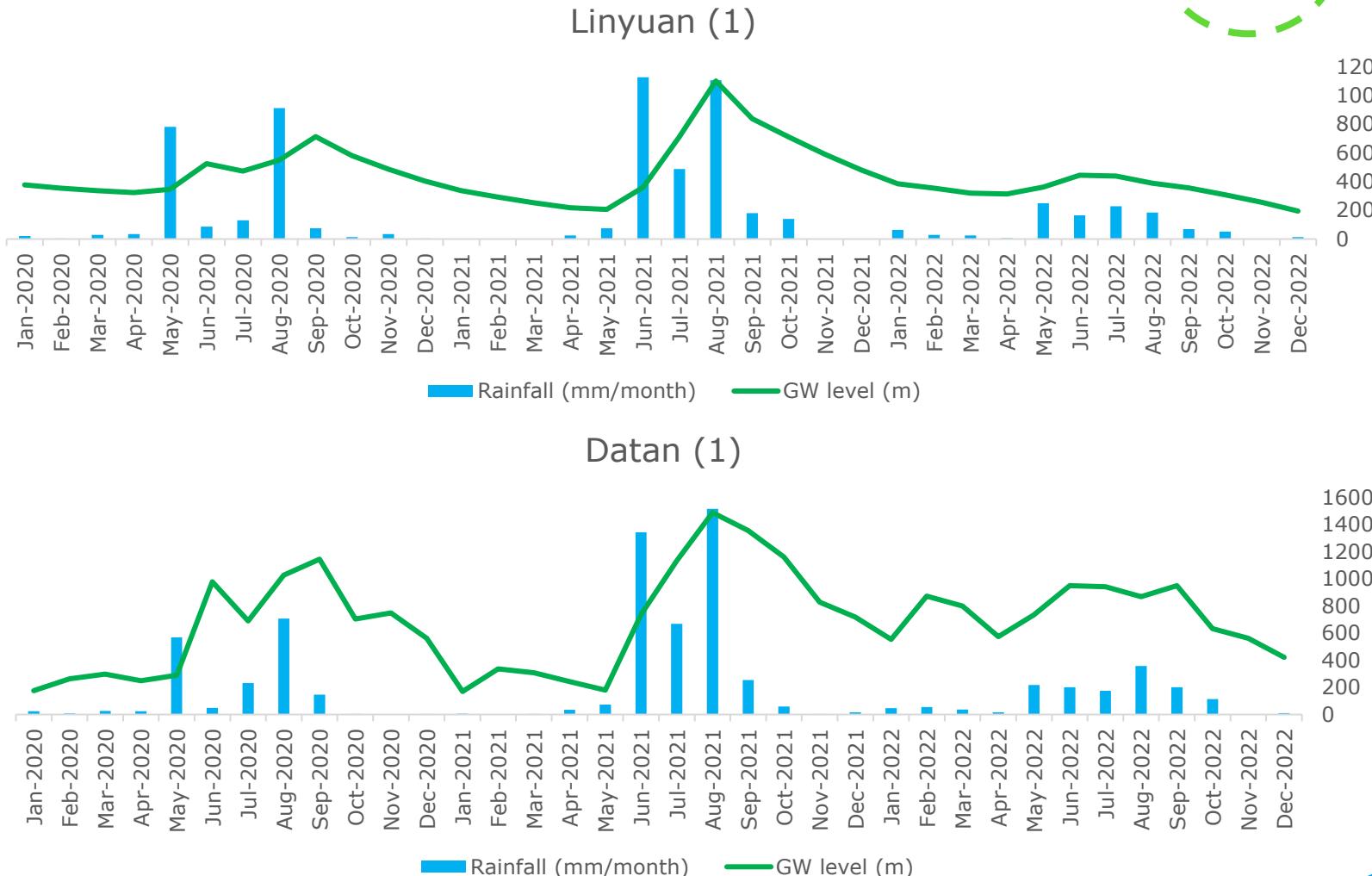
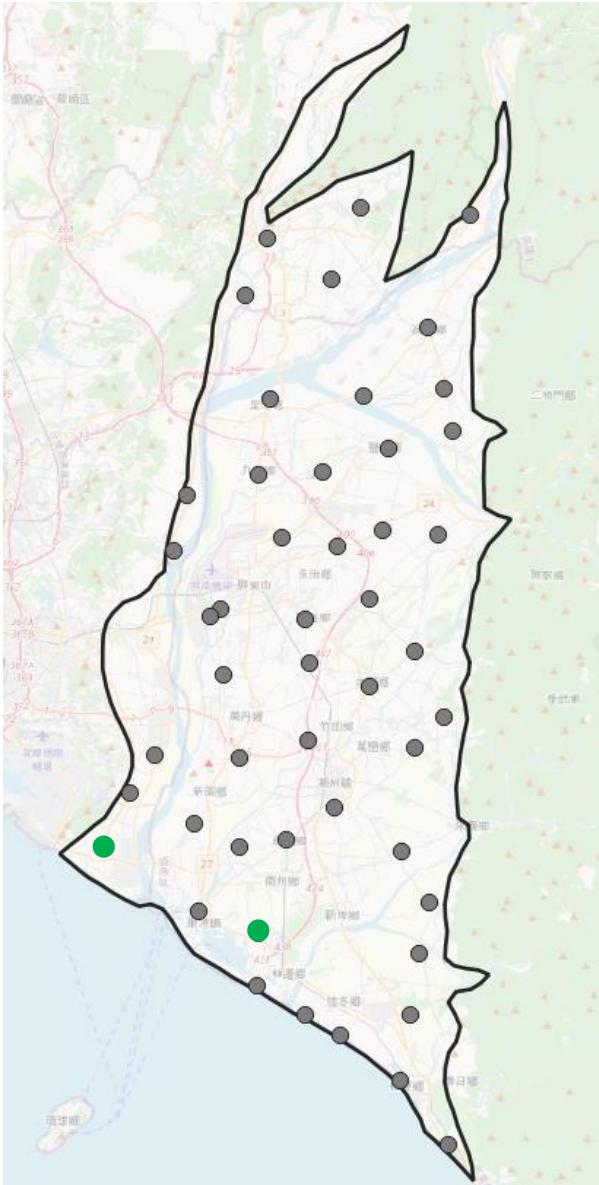
II

III

# PRELIMINARY RESULT

IV

## Groundwater level and rainfall distribution relationship



➤ Groundwater increase after a month from raining time

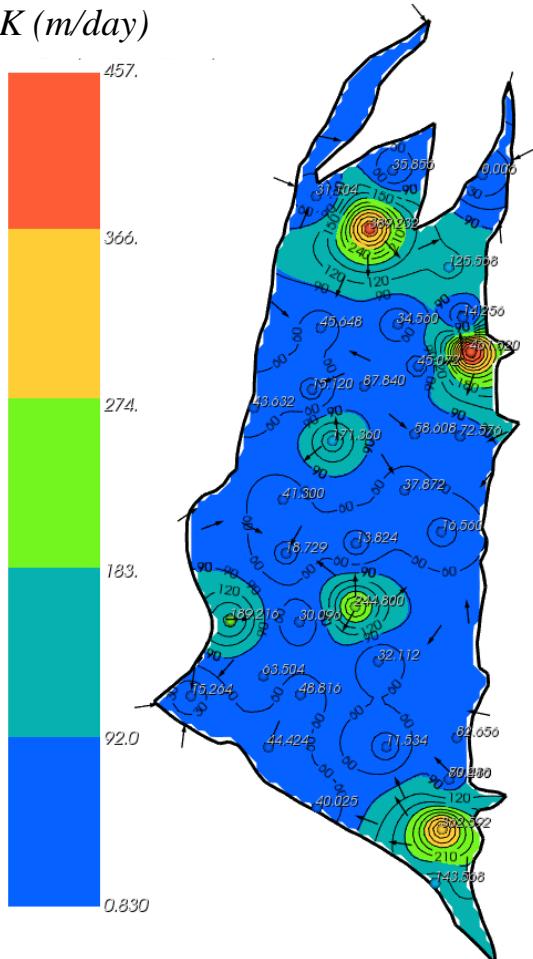
I

II

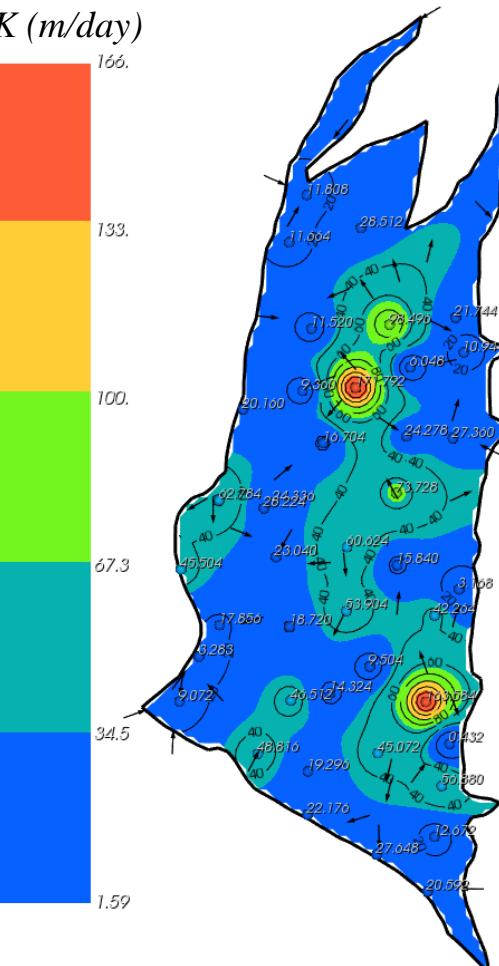
### PRELIMINARY RESULT

IV

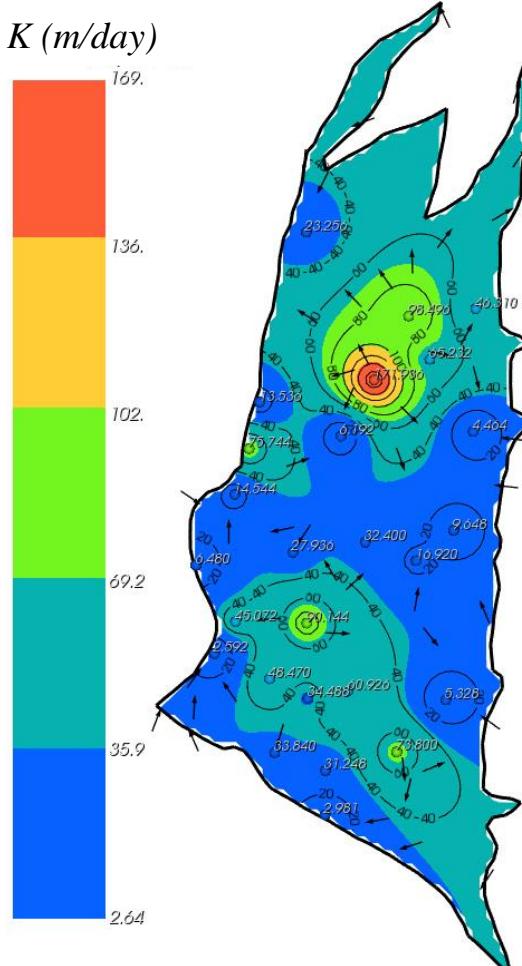
#### K value spatial distribution

 $K$  ( $m/day$ )

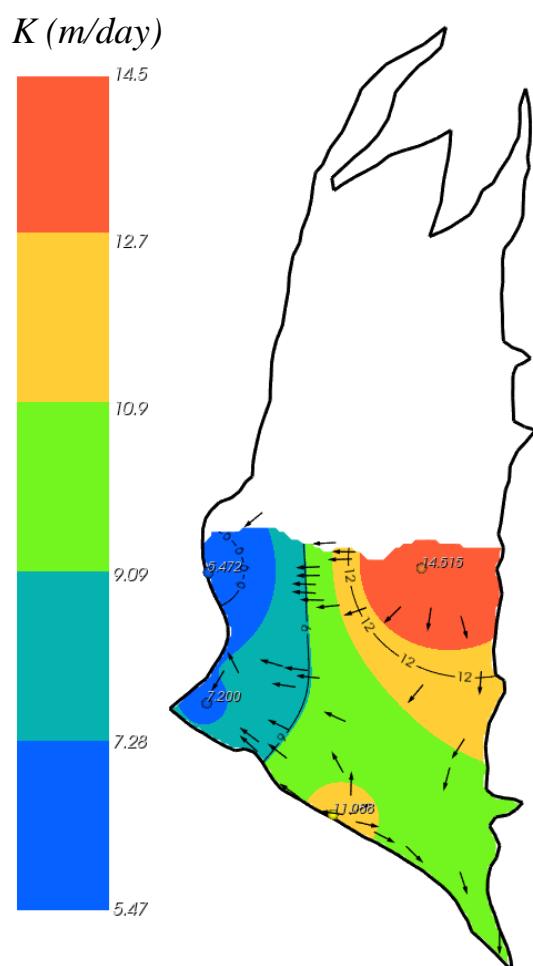
Aquifer 1

 $K$  ( $m/day$ )

Aquifer 2

 $K$  ( $m/day$ )

Aquifer 3

 $K$  ( $m/day$ )

Aquifer 4

I

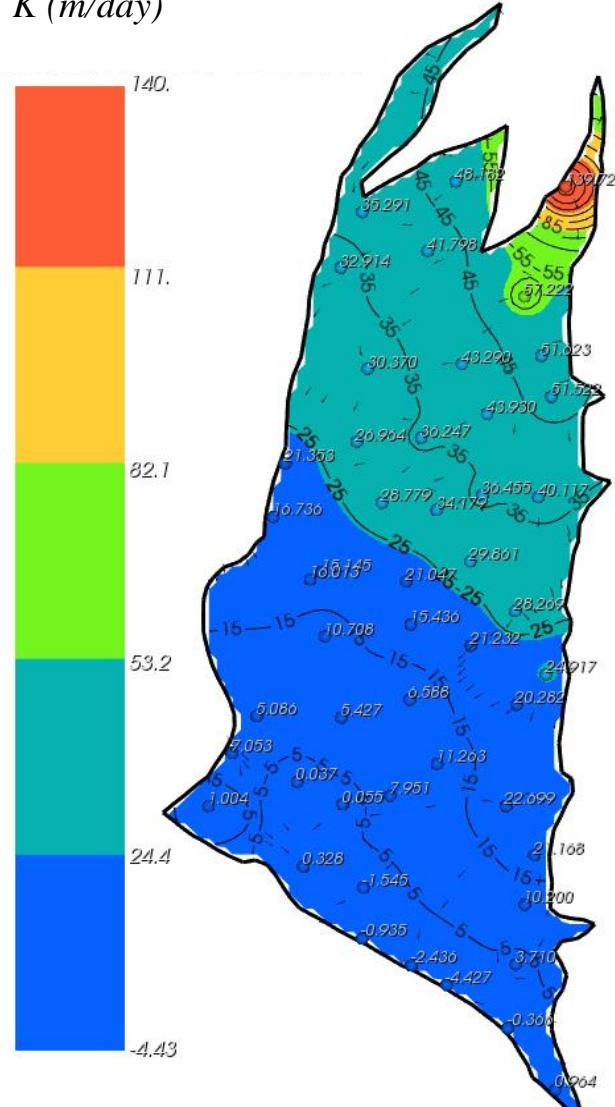
II

### III PRELIMINARY RESULT

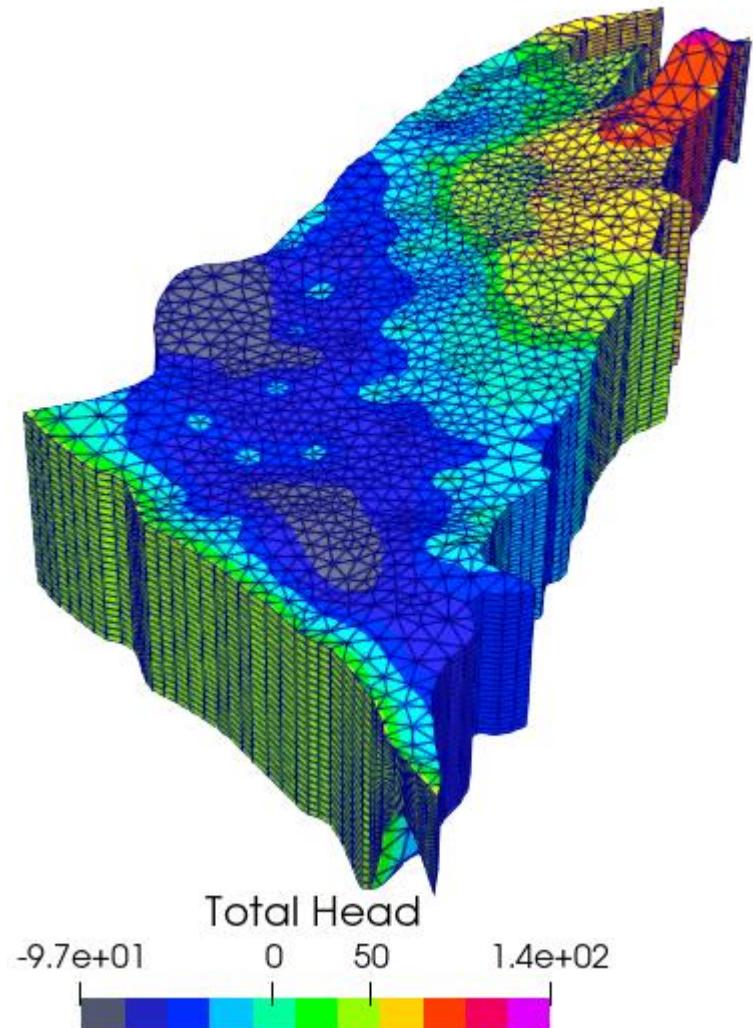
IV

## Steady-state simulation

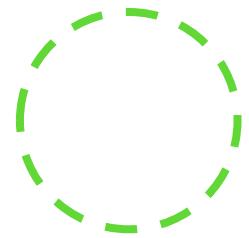
Rainfall average rate input:  $0,625 \times 10^{-3}$  m/day  
 $K$  (m/day)



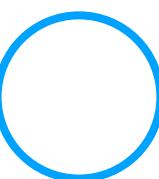
Initial result  
(2020 – Jan -01)



Initial head distribution results



## IV FUTURE WORK



1. Identify problem/define purpose

2. Build a conceptual model

3. Create mathematical models and select simulation programs

4. Model design  
From conceptual model to numerical model  
(Boundary and initial conditions/parameters/stress settings)

5. Calibration process

Selectivity calibrated

Calibrate model & evaluate accuracy

Calibration passed

Yes

Evaluate calibration results

No

6. Predict(Forecast Simulation)

7. Assessing uncertainty in predictive simulations

8. Evaluate results and prepare simulation reports

9. New field data collection completed  
Re-evaluate and refine the model

I

II

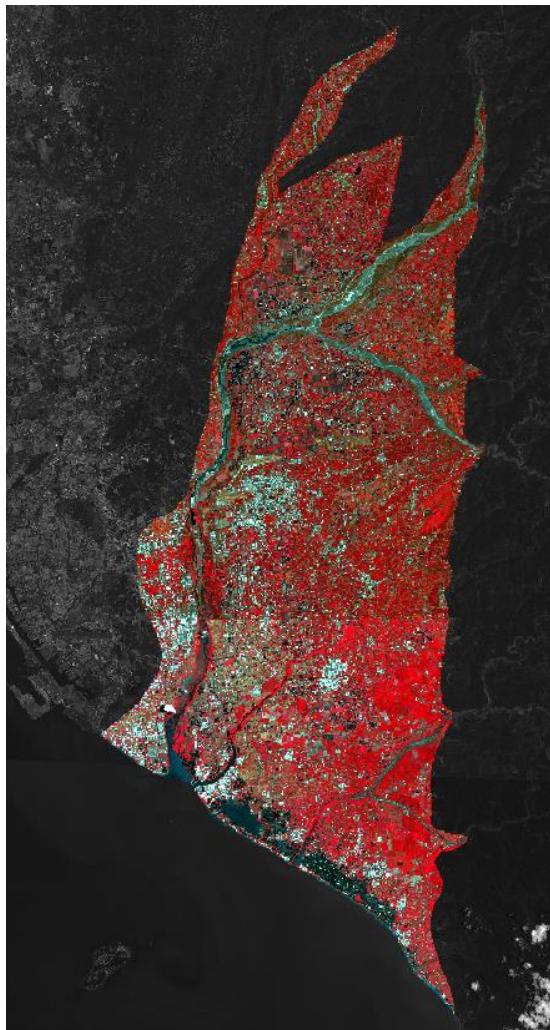
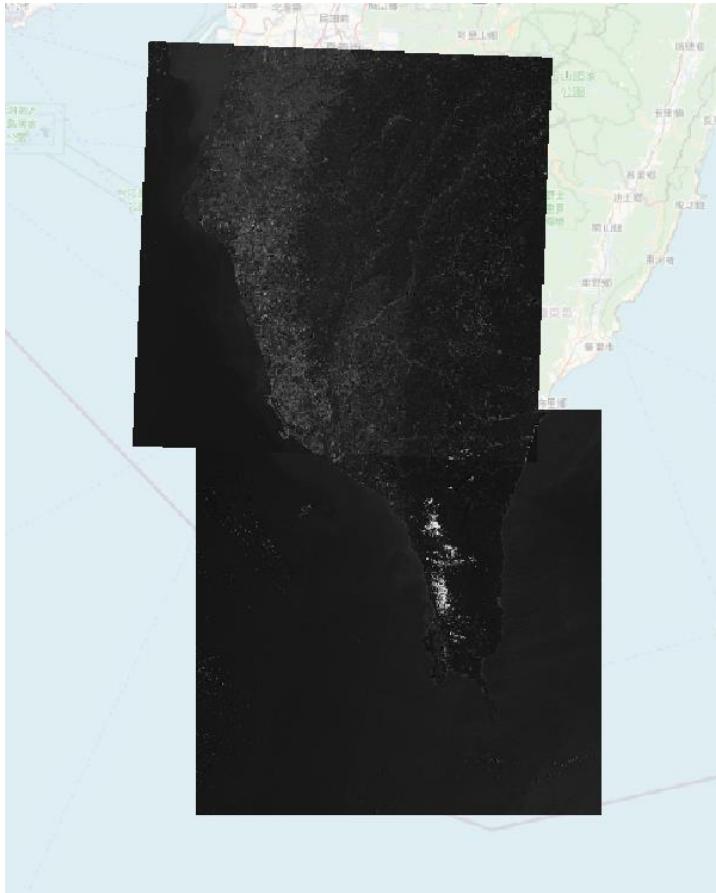
III

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## FUTURE WORK

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Using satellite image for classifying permeable and impermeable zone



*Mosaic with geometric correction*



*Target classification (source: USGS)*

THANK YOU  
FOR YOUR ATTENTION



Q & A



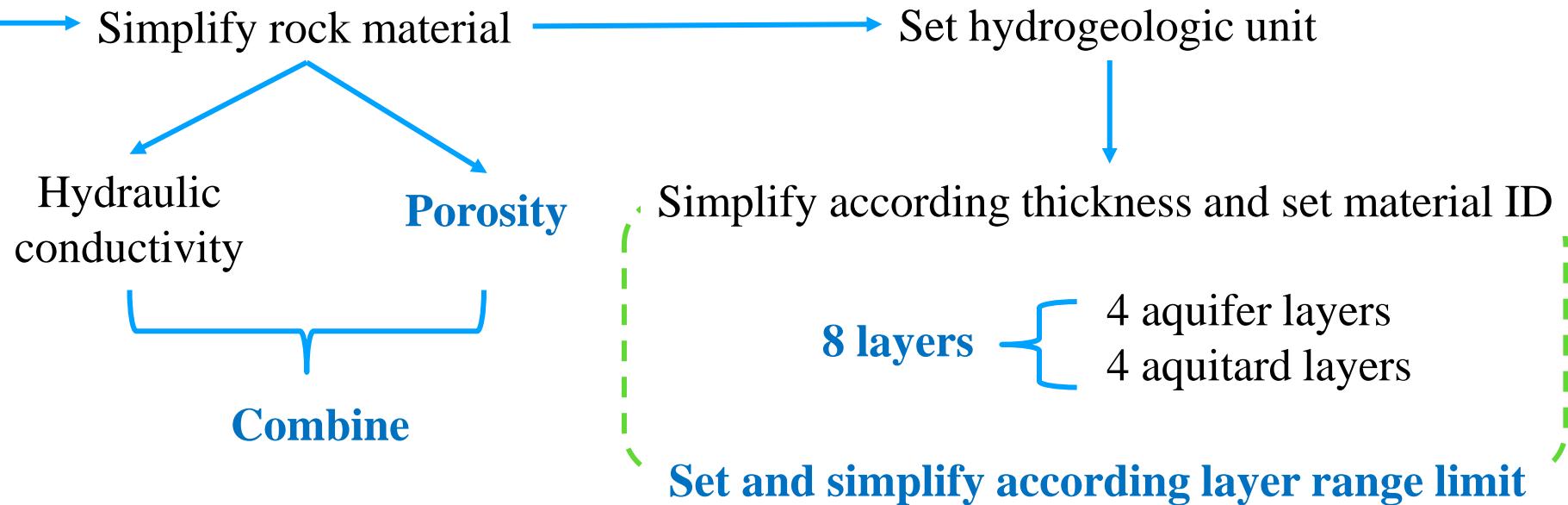
## FUTURE WORK

### Data collection (CGS)

Groundwater level

Borehole data  
(44 wells)

## SET UP STUDY DOMAIN - THMC

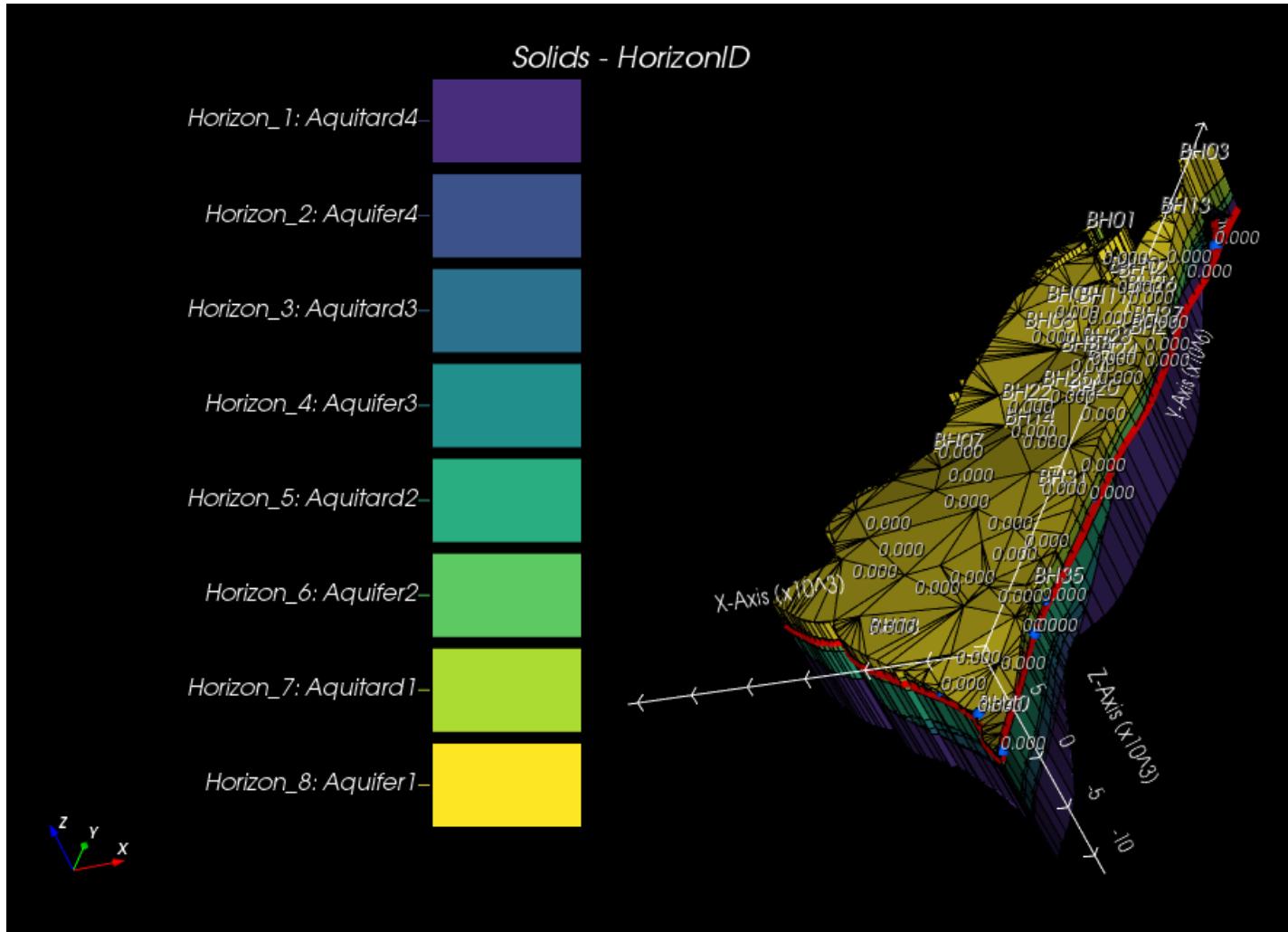
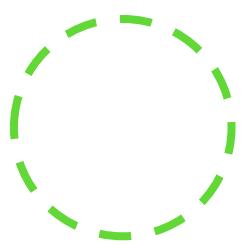




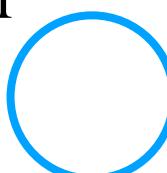
## FUTURE WORK



➤ Primarily simplified data test result for building study area domain



- The limited range of layers have not in expect
- Density in the middle too larger



I

## II METHODOLOGY

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Governing equation for flow through saturated-unsaturated media follow below equation:

$$\frac{\rho}{\rho_0} F \frac{\partial h}{\partial t} = \nabla \cdot \left[ K \cdot \left( \nabla h + \frac{p}{p_0} \nabla z \right) \right] + \frac{\rho^*}{\rho_0} q \quad (\text{Yeh et al., 1994a, 1994b})$$

$\theta$ : effective moisture content ( $L^3/L^3$ )

$h$ : pressure head (L)

$t$ : time (T)

$z$ : potential head (L)

$q$ : source/sink of fluid [ $(L^3/L^3)/T$ ]

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$\alpha'$ : modified compressibility of the soil matrix (1/L)

$\beta'$ : modified compressibility of the liquid (1/L)

$n_e$ : effective porosity ( $L^3/L^3$ )

$S$ : degree of effective saturation of water

$g$ : gravity ( $L/T^2$ )

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$K_{so}$ : referenced saturated hydraulic conductivity tensor (L/T)

$k_r$ : relative permeability or relative hydraulic conductivity (dimensionless)

$F$ : generalized storage coefficient (1/L)

$$F = \alpha' \frac{\theta}{n_e} + \beta' \theta + n_e \frac{dS}{dh}$$

$K$ : hydraulic conductivity tensor (L/T) with

$$K = \frac{\rho g}{\mu} k = \frac{\rho/\rho_0}{\mu/\mu_0} \frac{\rho_0 g}{\mu_0} k_s k_r = \frac{\rho/\rho_0}{\mu/\mu_0} K_{so} k_r$$

Darcy's velocity (L/T)

$$V = -K \cdot \left( \frac{\rho_0}{\rho} \nabla h + \nabla z \right)$$

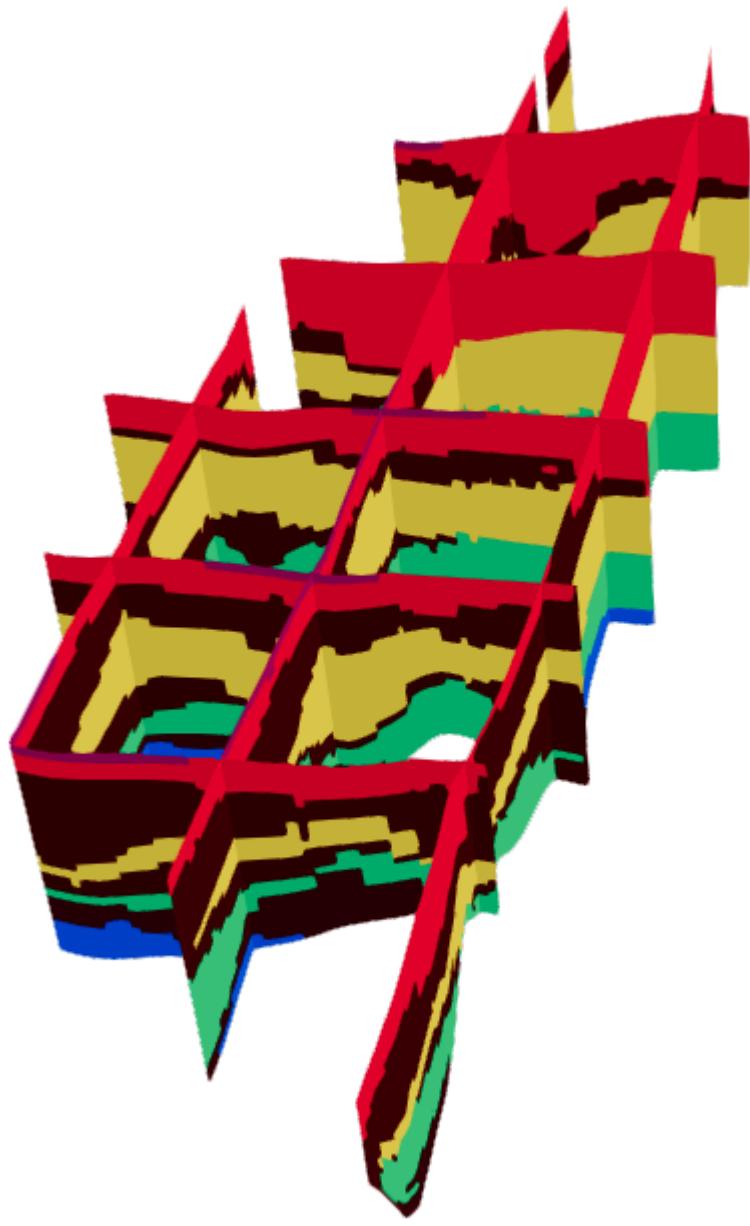
## Governing equation

- Conceptual hydrological: build up from 42 borehole cores (CGS website)
- K value reference:
  - Aquifer: 148 wells → Pumping test result from CGS report
  - Aquitard: According to type of rock and giving K value base on reference table of Domenico and Schwartz (1998)

Material	K lowest (m/day)	K highest (m/day)
Fine sand	0.017280000	17.280000
Silt	0.000086400	1.728000
Clay	0.000000864	0.000432

- Groundwater level: daily data from 2020 – 2022 of 122 stations (WRA)
- Rainfall rate: daily data from 2020 – 2022 of 8 stations (WRA)

## Data input



### Legendary

- Material 8
- Aquifer1
- Aquitard1
- Aquifer2
- Aquitard2
- Aquifer3
- Aquitard3
- Aquifer4

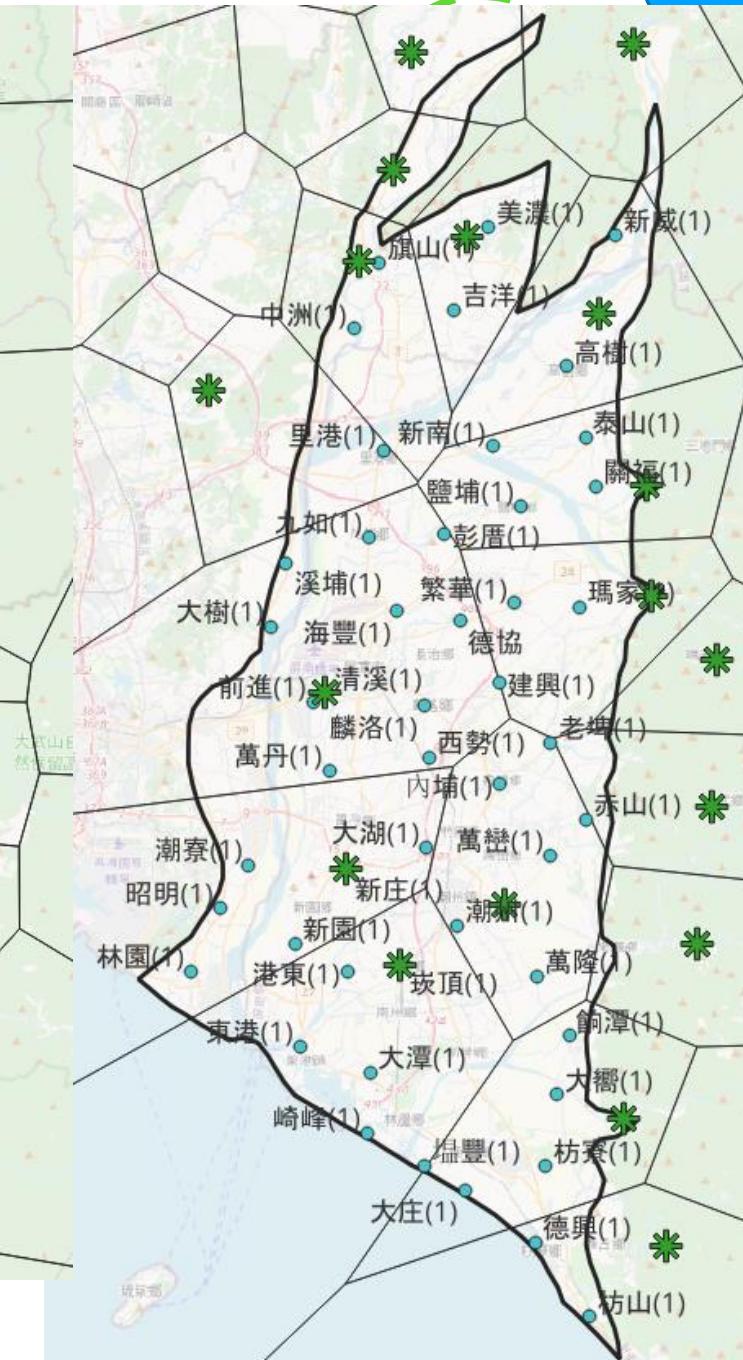
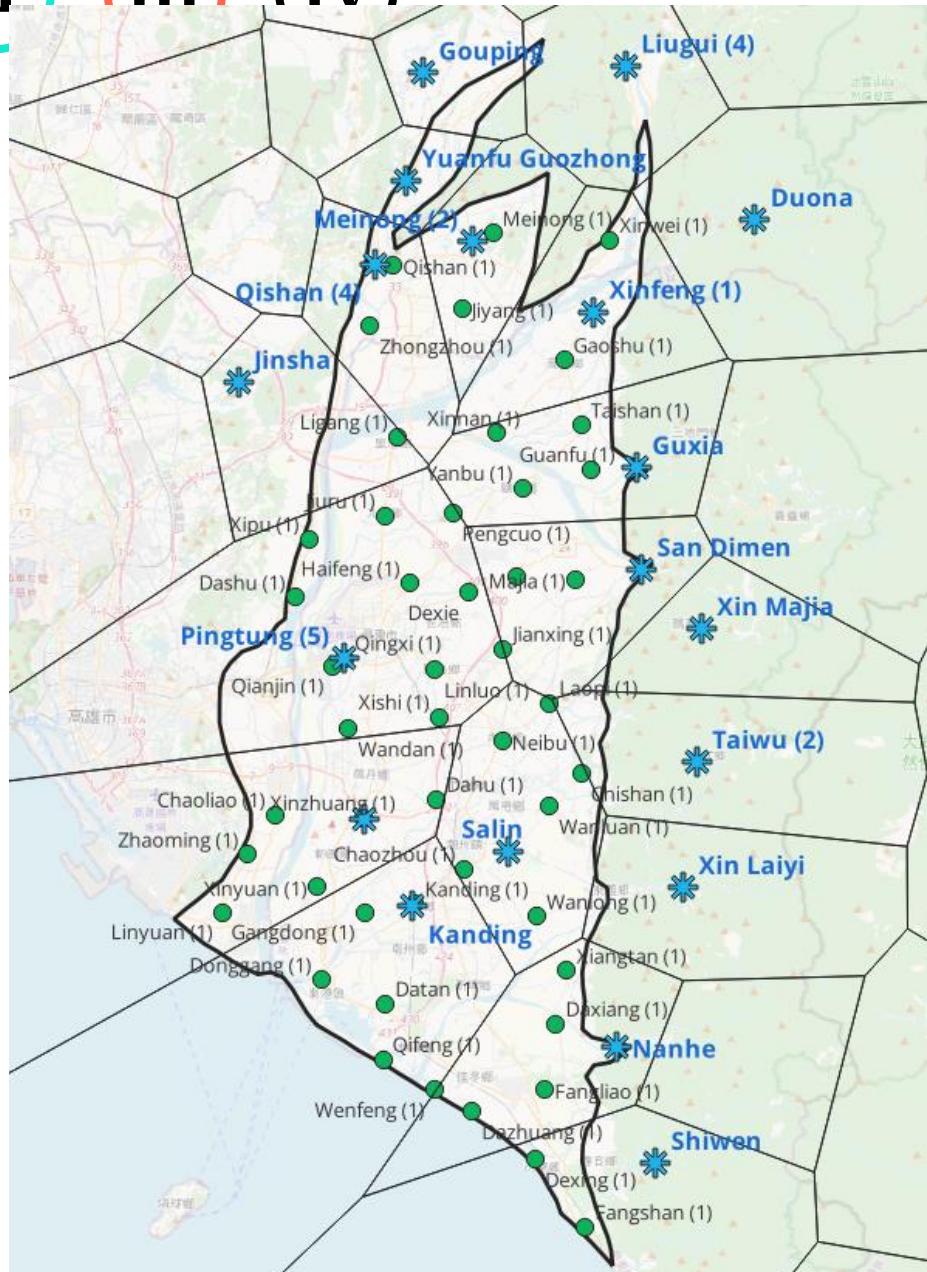
# I INTRODUCTION

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## THMC software



I

## II METHODOLOGY

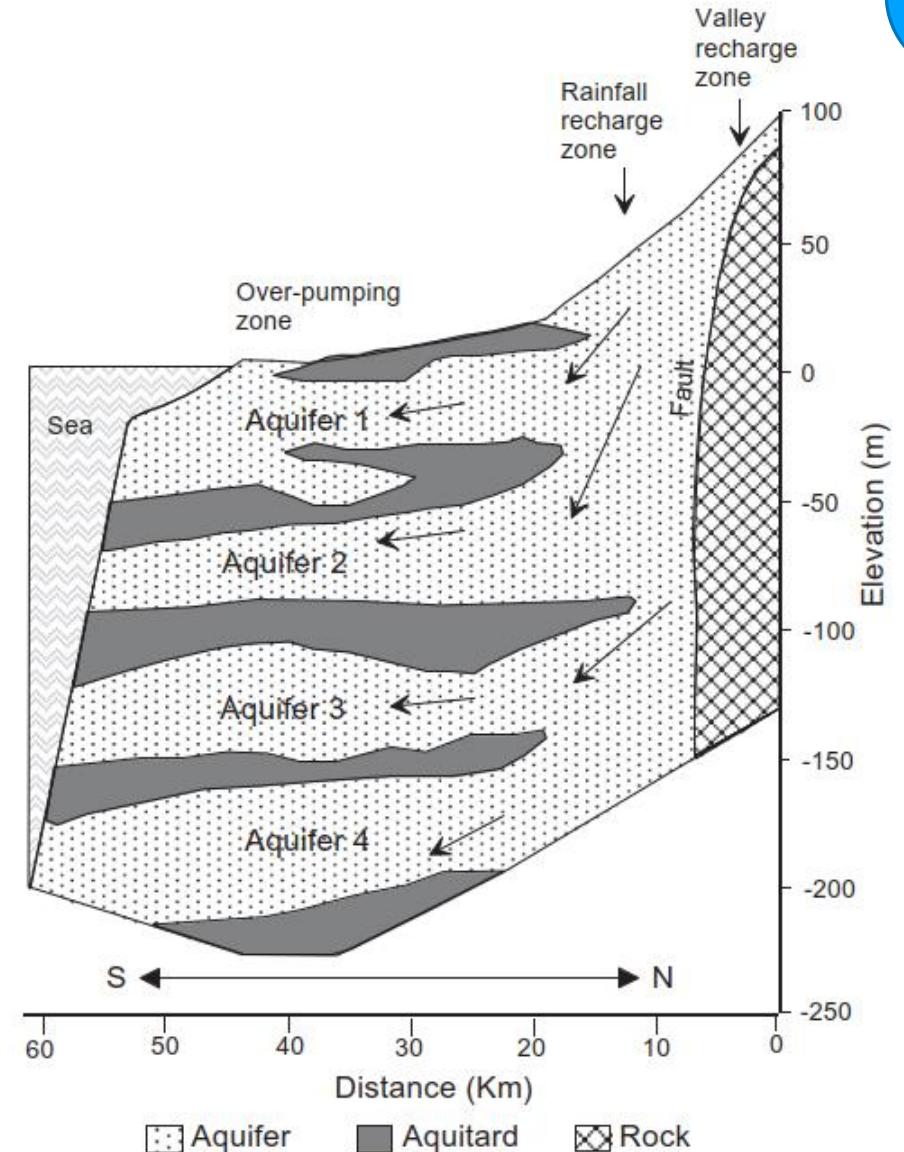
III

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• Hydrogeological  
 • Rainfall recharge zone  
 • Valley recharge zone

The plain sediments can divide into 8 layers:

- Aquifer: high permeable coarse sediments, ranging from medium sand to gravel.
- Aquitard: low permeable fine sediments, ranging from clay to fine sand



*A conceptual hydrogeological profile of the Pingtung Plain (Jang et al., 2016)*

I

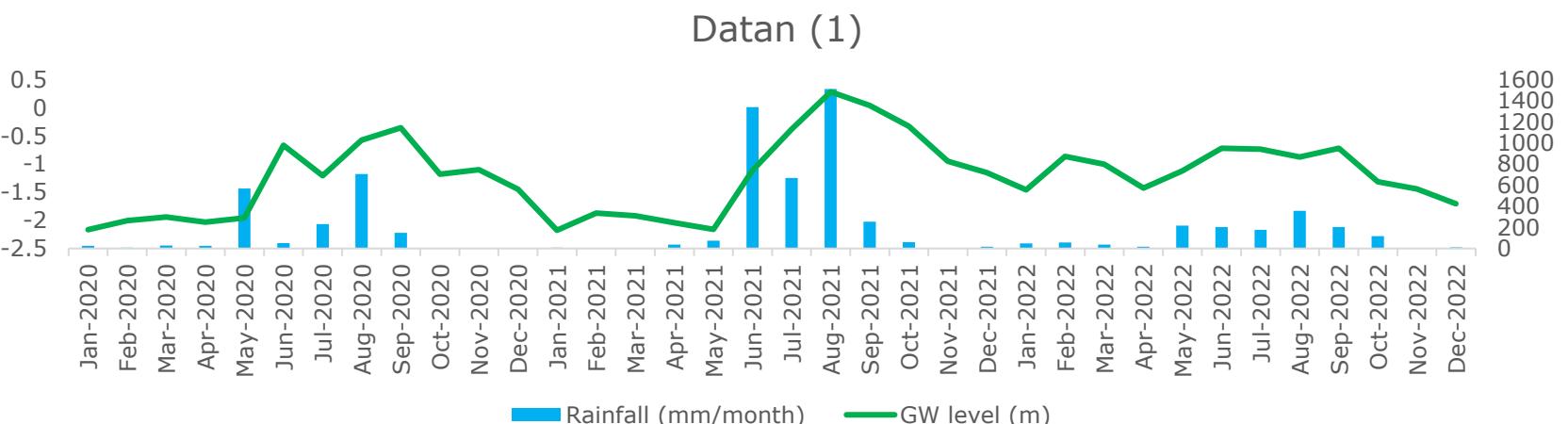
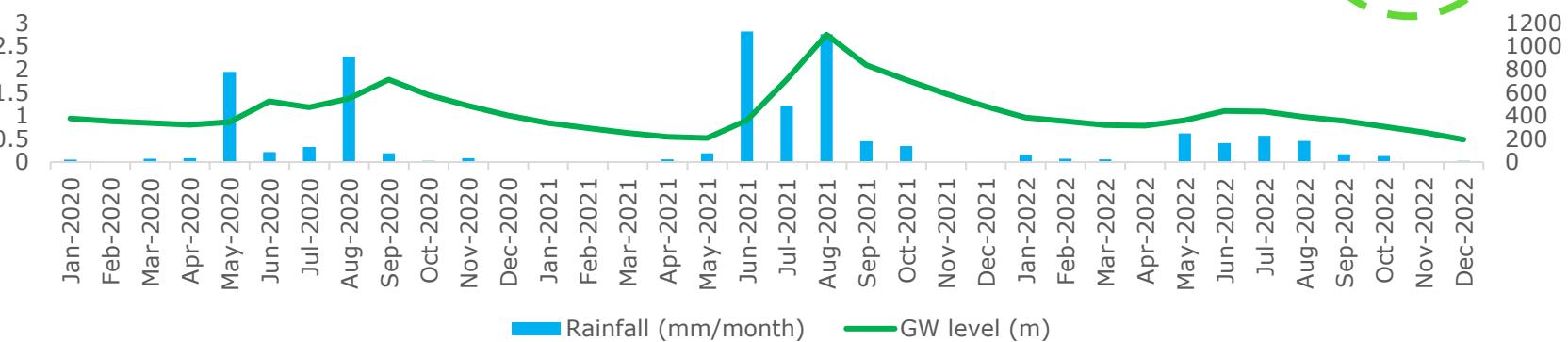
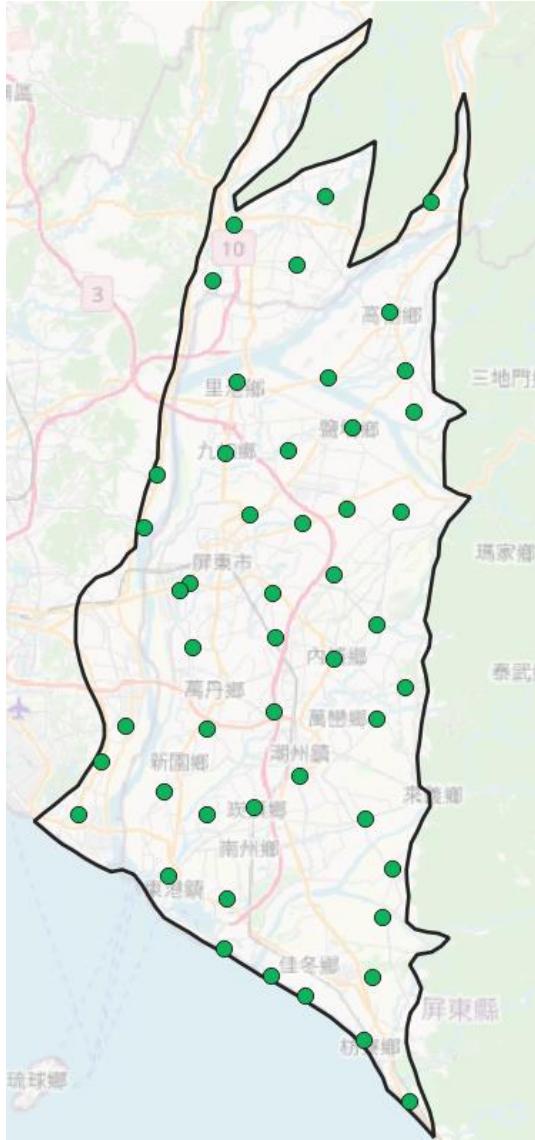
II

III

# PRELIMINARY RESULT

IV

## Groundwater level and rainfall distribution relationship



Groundwater increase after a month from raining time

I

II

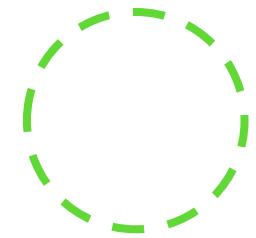
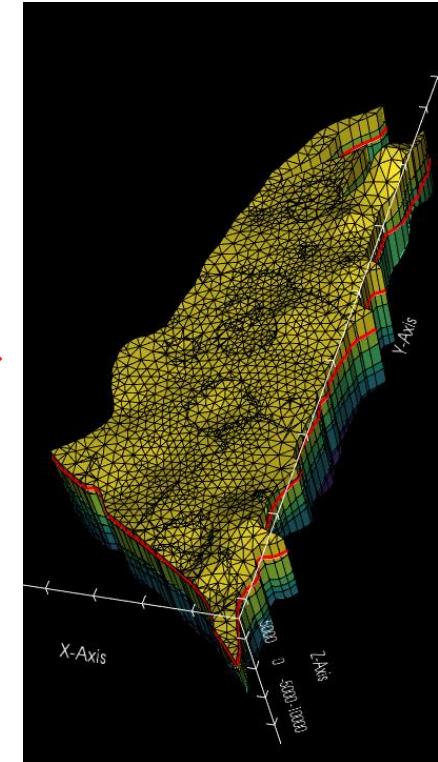
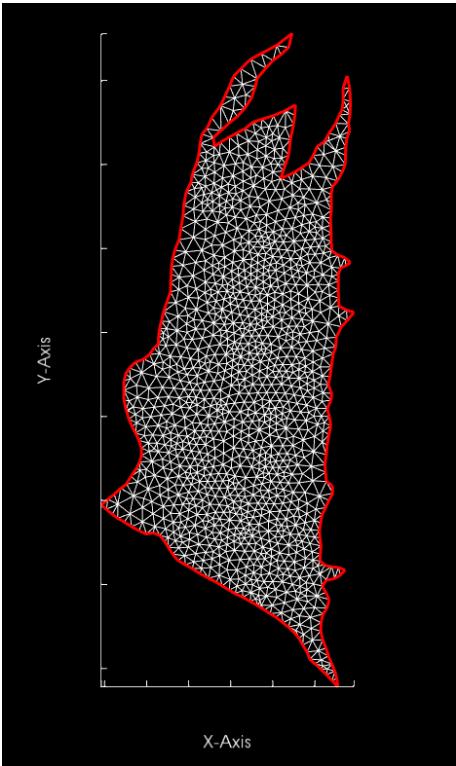
III

# PRELIMINARY RESULT

IV

## Conceptual hydrological domain

Create triangular grid and generate 3D domain



Uniform cells

